#### CRYSTALLOGRAPHY OF COPIAPITE

By

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THIS joint work is the result of interrupted studies which commenced in 1934 when Palache measured a number of crystals of copiapite from Chile and found them to be triclinic with a well developed form-series. In this part of the work the senior author received some assistance from Berman and Wolfe, and also from Peacock, who critically examined the chosen setting in relation to the general problem of choosing triclinic elements from goniometric data, and later made a direct determination of the crystal lattice from single crystal x-ray photographs which were kindly prepared by Buerger.

About this time Ungemach's important monograph on the sulphate minerals of Chile (1935) was received. In this detailed work, which includes descriptions of four new minerals and single-circle measurements on five triclinic species, Ungemach also discovered the triclinic symmetry of copiapite and described a most elaborate development of forms. However, it was found that Ungemach had chosen an entirely different setting from the one adopted at Harvard, and that his elements and angles do not agree closely with ours when transformed to our setting. Although Ungemach had given much thought to the setting of copiapite he generously conceded the propriety of the Harvard setting when the results of the x-ray measurements were known; however the differences in angles remained and no sufficient explanation for this was immediately apparent.

<sup>&</sup>lt;sup>1</sup>The results were transmitted personally to Ungemach in Strasbourg by Donnay, only a few days before Ungemach's untimely death on June 11, 1936. At that time Ungemach also bequeathed his large collection of measured crystals to Donnay, who in turn kindly placed the copiapite crystals at our disposal for further study.

In 1937-8, Berry undertook a study of the crystallography and crystal chemistry of copiapite during his first year of graduate study at Toronto. Peacock collaborated in this work which is described in an unpublished M.A. Thesis (Berry, 1938a), but only part of the work, that dealing with Ungemach's "pseudocopiapite" (Berry, 1938b), was published. Several abstracts of projected papers on copiapite and matters arising out of the study of that mineral have appeared: one announcing the "Harmonic-Arithmetic Rule" (Peacock, 1937); one giving the new geometrical and structural elements of copiapite (Peacock, 1939); and one on the variation of the composition and optical properties of the mineral (Berry, 1939).

During the past summer Palache returned to copiapite and, with Berry's Thesis and an angle-table by Peacock at hand, computed a complete angle-table for a projected volume of crystallographic tables, and a shorter presentation of the crystallography of the mineral for the second volume of the new Dana. When this was done it seemed proper that the previous scattered and interrupted work should be assembled in a single paper on the crystallography of this unusually interesting triclinic species. In the present account Palache is responsible for the measurements and calculations on the crystals from Chuquicamata (Table 1) and the formal angle-table (Table 4), while Peacock has prepared the rest of the paper with much assistance in the formal work by Berry. It is hoped that the results on the composition and optics of copiapite will be presented on another occasion.

#### EARLIER OBSERVATIONS

When distinctly crystallized, copiapite forms minute translucent yellow plates with perfect cleavage and pearly lustre on the plane of platy development. These plates are often rhombic in outline with principal optical directions practically coinciding with the normal to the plate and the diagonals of the rhomb. Copiapite was therefore considered to be orthorhombic by Bertrand (1881) and Des Cloizeaux (1881) and again by Posnjak & Merwin (1922).

On crystals showing small edge-faces Linck (1889) derived monoclinic elements:

 $a:b:c=0.4791:1:0.9759;\,\beta=108^{\circ}\,04'$ 

and a series of forms most of which have complex symbols. Linck's

elements are reproduced in Hintze (1892, p. 964) gave slightly different Goldschmidt (1913, p. 187); Schariz-Linck's measurements in an endeavou these recomputed values are given by

#### OBSERVATIONS ON CRYSTALS

Two-circle measurements. The tru apite was discovered independently a (Alsace), where Ungemach made a the material which had been studied by can be found between Linck's angles the present authors, but there is no crystallography is due to a misconcer apite, and that his results should be

The crystals measured in 1934 by Chuquicamata, Chile. They are smaplates, bevelled by several zones of n two are particularly strong; they hav of  $77^{\circ}$   $48\frac{1}{2}$ ' and they tend to give th appearance. Of these zones, the one was chosen as the (hk0)-zone, the other plane of flattening and perfect cleava could never be safely distinguished a preliminary measurement with  $(0\ 1938b,\ p.\ 11)$  preceded the regular adjustment on the vertical crystal ax

The composite plot from half a triclinic projection in which nearly all lie at the nodes of an oblique eccent which gives the a-axis shorter than the slope to the front and to the right of thangles gave the projection elements  $x_0' = 0.2156$ ,  $y_0' = 0.1078$ ,  $\nu = 79^{\circ}$  34' elements:

 $p_0: q_0: r_0 = 1.0065 : 0.4002 : 1; \lambda = 83$  $a:b:c = 0.4058 : 1 : 0.4039; \alpha = 93^c$  lertook a study of the crystallography and apite during his first year of graduate study llaborated in this work which is described in 'hesis (Berry, 1938a), but only part of the h Ungemach's "pseudocopiapite" (Berry

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#### RLIER OBSERVATIONS

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elements are reproduced in Hintze (1930, p. 4,415), while Dana (1892, p. 964) gave slightly different values which were adopted by  $G_{\rm oldschmidt}$  (1913, p. 187); Scharizer (1913, p. 384) recomputed Linck's measurements in an endeavour to simplify the symbols, and these recomputed values are given by Doelter (1929, p. 556).

#### OBSERVATIONS ON CRYSTALS FROM CHUQUICAMATA

Two-circle measurements. The true triclinic symmetry of copiapite was discovered independently at Harvard and at Strasbourg (Alsace), where Ungemach made a thorough re-examination of the material which had been studied by Linck. Some correspondence can be found between Linck's angles and those of Ungemach and the present authors, but there is no doubt that Linck's unnatural crystallography is due to a misconception of the symmetry of copiapite, and that his results should be discarded as erroneous.

The crystals measured in 1934 by Palache were obtained from Chuquicamata, Chile. They are small, often somewhat elongated plates, bevelled by several zones of narrow facets. Of these zones two are particularly strong; they have an average interzonal angle of 77°  $48\frac{1}{2}$ ′ and they tend to give the plates their pseudorhombic appearance. Of these zones, the one with the generally longer edge was chosen as the (hk0)-zone, the other as the (0kl)-zone, with the plane of flattening and perfect cleavage as (010). These two zones could never be safely distinguished by inspection, and therefore a preliminary measurement with (010) as the pole-face (Berry, 1938b, p. 11) preceded the regular two-circle measurement with adjustment on the vertical crystal axis.

The composite plot from half a dozen crystals gave a typical triclinic projection in which nearly all the poles of the terminal faces lie at the nodes of an oblique eccentric net. In the one position which gives the *a*-axis shorter than the *b*-axis and makes the *c*-plane slope to the front and to the right of the *a*-plane, the mean measured angles gave the projection elements:  $p_0' = 1.0353$ ,  $q_0' = 0.4117$ ;  $x_0' = 0.2156$ ,  $y_0' = 0.1078$ ,  $v = 79^\circ$  34' and hence the polar and linear elements

 $p_0: q_0: r_0 = 1.0065: 0.4002: 1; \lambda = 83^{\circ} 59', \mu = 76^{\circ} 59\frac{1}{2}', \nu = 79^{\circ} 34'$   $a:b:c = 0.4058: 1: 0.4039; \alpha = 93^{\circ} 50', \beta = 102^{\circ} 10', \gamma = 99^{\circ} 21\frac{1}{2}'$ 

TABLE 1 COPIAPITE: MEASURED AND CALCULATED ANGLES ON CRYSTALS FROM CHUOUICAMATA

F	Mea	sured		C	alculated			01
Forms	arphi	ρ	φ	ρ	A	В	С	Obs
(001)	63°11′	13°36′	63°26′	13°33′	76°59½	′ 83°59′	0°00′	5
(010)	0 00	90 00	0 00	90 00	<b>7</b> 9 <b>3</b> 4	0 00	83 59	14
(100)	79 34	90 00	79 34	90 00	0 00	79 34	$76\ 59\frac{1}{2}$	10
(110)	59 41	90 00	59 31½	90 00	$20\ 02\frac{1}{2}$		<b>7</b> 6 <b>2</b> 9	3
$(1\overline{1}0)$	102 33	90 00	102 25	90 00	22 51	102 25	$79\ 30\frac{1}{2}$	
$(2\overline{3}0)$	112 56	90 00	$112 \ 53\frac{1}{2}$	90 00	$33\ 19\frac{1}{2}$	$112 \ 53\frac{1}{2}$	$81\ 14\frac{1}{2}$	1
$(1\bar{2}0)$	122 08	90 00	121 59	90 00	42 25	121 59	$82\ 58\frac{1}{2}$	
$(2\overline{5}0)$	129 45	90 00	129 34 ½	90 00		$129\ 34\frac{1}{2}$	$84\ 33\frac{1}{2}$	
$(1\overline{3}0)$	135 56	90 00	135 49	90 00	$55\ 35\frac{1}{2}$	135 49	85 55	2
$(1\bar{7}0)$	159 42	90 00	159 18	90 00	<b>7</b> 9 <b>4</b> 4	159 18	$91\ 22\frac{1}{2}$	2
(011)	22 35	<b>29 17</b>	$22\ 32\frac{1}{2}$	$29\ 21\frac{1}{2}$			$20\ 54\frac{1}{2}$	
(021)	13 08	43 09	13 02	$43\ 42\frac{1}{2}$	$74\ 01\frac{1}{2}$	$47 \ 41\frac{1}{2}$	$36\ 17\frac{1}{2}$	1
$(0\overline{1}1)$	144 28	20 36	144 39	20 26	$81\ 32\frac{1}{2}$	$106\ 32\frac{1}{2}$	$22\ 33\frac{1}{2}$	2
$(0\overline{2}1)$	162 40	36 46	163 14	$36\ 46\frac{1}{2}$	86 13	$124\ 58\frac{1}{2}$	$40\ 59\frac{1}{2}$	1
$(0\overline{3}1)$	169 33	49 37	$169 \ 10\frac{1}{2}$	48 56	$89\ 41\frac{1}{2}$	$137 \ 46\frac{1}{2}$	53 47½	1
(101)	<b>7</b> 6 15	$\frac{1}{2}$ 51 51 $\frac{1}{2}$	76 32½	51 45	38 21	<b>7</b> 9 <b>2</b> 8	$38\ 38\frac{1}{2}$	2
$(\overline{1}02)$	$-86 \ 11$	16 04	$-87\ 15$	-	105 56	$89\ 13\frac{1}{2}$	$28\ 56\frac{1}{2}$	
$(\overline{1}01)$	<b>-96 08</b>	$38\ 47\frac{1}{2}$	$-95 \ 40\frac{1}{2}$	$38\ 53\frac{1}{2}$	128 44	$93\ 33\frac{1}{2}$	$51\ 44\frac{1}{2}$	6
$(\overline{2}01)$	-97 36	61 22	-98 21	61 29	$151\ 24\frac{1}{2}$	97 20	<b>7</b> 4 25	3
$(\overline{1}11)$	-67 27	40 53	$-67\ 31\frac{1}{2}$	$40\ 58\frac{1}{2}$	123 24	<b>75 29</b>	$50\ 42\frac{1}{2}$	
$(\overline{1}\overline{1}1)$	-121 56	43 43	$-121\ 28\frac{1}{2}$	$43\ 15\frac{1}{2}$	129 46	110 58	$56\ 46\frac{1}{2}$	2
$(\overline{1}32)$	-26 22	34 18	$-24\ 55\frac{1}{2}$	$34\ 51\frac{1}{2}$	$98\ 13\frac{1}{2}$	58 47	36 43	1
(121)	44 54	60 18	47 48	59 01	$43\ 12\frac{1}{2}$	$54\ 50\frac{1}{2}$	$46\ 03\frac{1}{2}$	1
(131)	39 13	62 56	$38 \ 52\frac{1}{2}$	$63\ 02\frac{1}{2}$	<b>47 2</b> 9	$46\ 03\frac{1}{2}$	$50\ 53\frac{1}{2}$	1
$(\overline{1}31)$	-34 22	54 55	-34 47	54 35½	109 38½	$47 58\frac{1}{2}$	57 35½	
$(\overline{2}11)$	-87 05	61 06	-85 28	61 18	$147 \ 55\frac{1}{2}$		$73\ 05\frac{1}{2}$	
$(\overline{4}52)$	-68 16	62 52	$-67\ 17\frac{1}{2}$	63 08	$138 \ 19\frac{1}{2}$	$69 \ 51\frac{1}{2}$	72 21	1
$(\overline{2}31)$	-63 15	64 25	$-62\ 00\frac{1}{2}$	$64\ 07\frac{1}{2}$	134 49	65 01	$72\ 25\frac{1}{2}$	
$(\bar{4}72)$	-5728	64 58	$-57 \ 11\frac{1}{2}$	$65\ 13\frac{1}{2}$	$131\ 24\frac{1}{2}$	60 32	72 26	2

CRYSTALLOGRAPHY OF CO

The mean measured angles are summ lated from these elements in Table 1.

Fig. 1 shows a crystal of copiapite fron tinct elongation in the direction of the somewhat irregular development of the be was sketched and measured by Berman used for the x-ray measurements by Peac tive of the material which was analysed b given later. Other crystals showed furthe habit with occasional approach to the pse crystal from Sierra Gorda shown in Fig. 3

*X-ray measurements*. The crystal sho as for two-circle measurement, to rotate a cobalt radiation, a rotation photograph v measurement of the c-period; a Weissenberg layer line gave the principal spacings, d(angle  $\nu$  (or  $\gamma^*$ ) included by the normals to a Weissenberg resolution of the first layerlayer of the reciprocal lattice from which  $\alpha^*$ ) and  $\mu$  (or  $\beta^*$ ) were calculated. By met are now standard practice these quantitie crystal lattice cell:

$$a = 7.33$$
,  $b = 18.15$ ,  $c = 7.27$  kX;  $\alpha = 93°51'$ 

The reciprocal lattice projection of t metrically similar to the gnomonic net of procedure of choosing the single position c naming its elements is strictly analogous t sing the geometrical projection. It can edges of the chosen lattice cell are the th periods of the lattice, and that this cell is s single setting which conforms to the ru obtuse.

The geometrical and structural axial Chuquicamata compare as follows:

> a:b:c0.4058 : 1 : 0.4039 93°50′ 102°10′ 99 0.4037:1:0.4005 93°51' 101°30' 99

TABLE 1
\*\*ASURED AND CALCULATED ANGLES
\*\*TALS FROM CHUQUICAMATA

	Ca	lculated			Obs.
φ	ρ	A	В	С	
63°26′	13° 33′	76°59½′	83°59′	0°00′	5
0 00	90 00	<b>7</b> 9 <b>3</b> 4	0 00	83 59	14
<b>7</b> 9 <b>3</b> 4	90 00	0 00	<b>7</b> 9 34	76 59½	10
59 31½	90 00	$20\ 02\frac{1}{2}$	$59\ 31\frac{1}{2}$		3
102 25	90 00	22 51		$79\ 30\frac{1}{2}$	8
$112 \ 53\frac{1}{2}$	90 00	$33 \ 19\frac{1}{2}$	$112 \ 53\frac{1}{2}$	$81 \ 14\frac{1}{2}$	1
121 59	90 00		121 59	$82\ 58\frac{1}{2}$	10
$129 \ 34\frac{1}{2}$	90 00		$129\ 34\frac{1}{2}$	$84 \ 33\frac{1}{2}$	1
135 49	90 00	$55\ 35\frac{1}{2}$	135 49	85 55	2
159 18	90 00	79 44	159 18	$91\ 22\frac{1}{2}$	2
$22\ 32\frac{1}{2}$	$29\ 21\frac{1}{2}$		$63\ 04\frac{1}{2}$	$20\ 54\frac{1}{2}$	3
13 02	$43\ 42\frac{1}{2}$	$74\ 01\frac{1}{2}$	$47\ 41\frac{1}{2}$	$36\ 17\frac{1}{2}$	1
144 39	20 26	81 32½	$106\ 32\frac{1}{2}$	$22\ 33\frac{1}{2}$	2
163 14	$36\ 46\frac{1}{2}$	86 13	$124\ 58\frac{1}{2}$		
$169 \ 10^{\frac{1}{2}}$	48 56	$89\ 41\frac{1}{2}$	$137 \ 46\frac{1}{2}$	$53\ 47\frac{1}{2}$	1
$76\ 32\frac{1}{2}$	51 45	38 21	<b>7</b> 9 <b>2</b> 8	$38\ 38\frac{1}{2}$	2
-87 15		105 56	$89\ 13\frac{1}{2}$	$28\ 56\frac{1}{2}$	
$-95\ 40\frac{1}{2}$	$38 \ 53\frac{1}{2}$	128 44	$93\ 33\frac{1}{2}$	$51\ 44\frac{1}{2}$	6
-98 21	61,29	$151\ 24\frac{1}{2}$	97 20	<b>74</b> 25	3
$-67\ 31\frac{1}{2}$		123 24	<b>75 29</b>	$50\ 42\frac{1}{2}$	
$\cdot 121 \ 28\frac{1}{2}$	$43\ 15\frac{1}{2}$	129 46	110 58	$56\ 46\frac{1}{2}$	2
$-24\ 55\frac{1}{2}$	34 51½	98 13½	58 47	36 43	1
47 48			$54 \ 50\frac{1}{2}$	$46\ 03\frac{1}{2}$	1
$38\ 52\frac{1}{2}$		<b>47 2</b> 9		$50 \ 53\frac{1}{2}$	1
-34 47	54 35½	109 38	½ 47 58½	57 35½	1
$-85\ 28$	61 18			73 05	
$-67  17\frac{1}{2}$				<b>72</b> 21	1
$-62\ 00^{\frac{1}{2}}$	64 07	134 49	65 01	<b>72</b> 25	
$-57 \ 11\frac{1}{2}$	65 13	131 24	60 32	<b>72</b> 26	2

The mean measured angles are summarized with angles calculated from these elements in Table 1.

Fig. 1 shows a crystal of copiapite from Chuquicamata with distinct elongation in the direction of the chosen vertical axis and somewhat irregular development of the bevelling zones. This crystal was sketched and measured by Berman and it was subsequently used for the x-ray measurements by Peacock; it is also representative of the material which was analysed by Gonyer with the results given later. Other crystals showed further variations of the tabular habit with occasional approach to the pseudorhombic habit of the crystal from Sierra Gorda shown in Fig. 3.

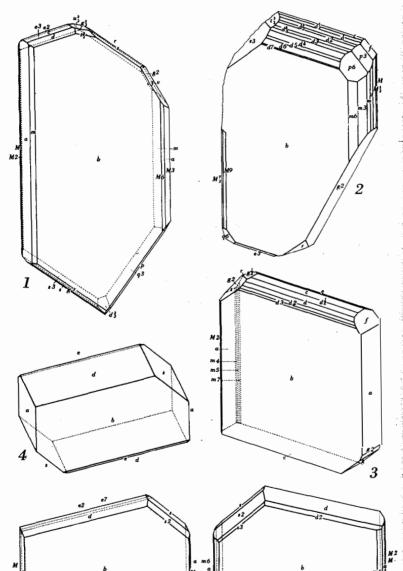
X-ray measurements. The crystal shown in Fig. 1 was adjusted, as for two-circle measurement, to rotate about the long axis. Using cobalt radiation, a rotation photograph was taken giving a direct measurement of the c-period; a Weissenberg resolution of the zero-layer line gave the principal spacings, d(100) and d(010), and the angle  $\nu$  (or  $\gamma^*$ ) included by the normals to these sets of planes; and a Weissenberg resolution of the first layer-line gave the offset of this layer of the reciprocal lattice from which the remaining angles  $\lambda$  (or  $a^*$ ) and  $\mu$  (or  $\beta^*$ ) were calculated. By methods of calculation which are now standard practice these quantities give the elements of the crystal lattice cell:

$$a = 7.33$$
,  $b = 18.15$ ,  $c = 7.27$  kX;  $\alpha = 93^{\circ}51'$ ,  $\beta = 101^{\circ}30'$ ,  $\gamma = 99^{\circ}23'$ 

The reciprocal lattice projection of the first layer-line is geometrically similar to the gnomonic net of the (hk1) planes, and the procedure of choosing the single position of the x-ray projection and naming its elements is strictly analogous to the procedure in discussing the geometrical projection. It can easily be shown that the edges of the chosen lattice cell are the three shortest non-coplanar periods of the lattice, and that this cell is set in the usually preferred single setting which conforms to the rules: c < a < b; a and b obtuse.

The geometrical and structural axial ratios of copiapite from Chuquicamata compare as follows:

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a:b:c \alpha \beta \gamma 0.4058:1:0.4039 93°50' 102°10' 99°21\frac{1}{2}' (Gon., Palache) 0.4037:1:0.4005 93°51' 101°30' 99°23' (X-ray, Peacock)
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CONTRIBUTIONS TO CANADIAN MINERALOGY

The substantial agreement shows that the tural methods led to one and the same cry *Composition*. On this material Berm gravity 2.154 and Gonyer made the analysis.

TABLE 2
COPIAPITE: ANALYSIS (GONYER) AN

Analy	/sis		Atoms in U				
FeO Al <sub>2</sub> O <sub>3</sub> Fe <sub>2</sub> O <sub>3</sub> SO <sub>3</sub> H <sub>2</sub> O	0.44 1.72 27.28 39.83 29.92	Fe" Al Fe" S H <sub>2</sub> O	0.08 0.41 4.18 6.09 20.34	Oxyge			
Insol.	0.55						

99.74

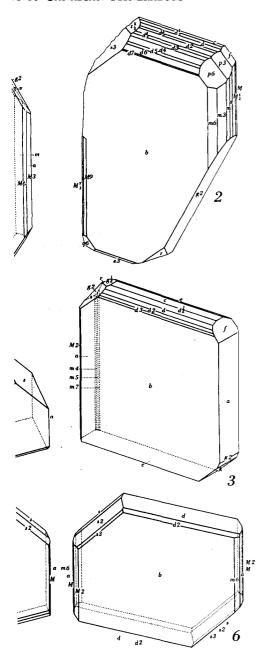
The atomic content approaches X(OH) X is one oxygen equivalent of (Fe''', Al) a fuller discussion of this and other analyse for the projected paper on the composition

### OBSERVATIONS ON CRYSTALS FRO

Ungemach's extensive work on copia complex crystals from Sierra Gorda. At with the plane of platy development a ("orientation ancienne") he finally chang this plane is the side pinakoid ("notation spect Ungemach's final setting resembles

Figs. 1-6.—Crystals of copiapite from Ch material used for 2-circle measurement, x-ray meas 2, 3.—Sierra Gorda; two of Ungemach's crystals r position. Figs. 4, 5, 6.—Tierra Amarilla (pseudoc crystals remeasured and redrawn in new position end of the c-axis is directed upward and the crystal of b(010); in Figs. 2, 3, 6, representing crystals whinegative end of the c-axis, this end is directed upward from the direction of  $b(0\overline{10})$ .

IS TO CANADIAN MINERALOGY



The substantial agreement shows that the morphological and structural methods led to one and the same crystal lattice cell.

Composition. On this material Berman measured the specific gravity 2.154 and Gonyer made the analysis given in Table 2.

TABLE 2
COPIAPITE: ANALYSIS (GONYER) AND CELL CONTENT

Analy	ysis		Atoms in Unit Cell				
FeO Al <sub>2</sub> O <sub>3</sub> Fe <sub>2</sub> O <sub>3</sub> SO <sub>3</sub> H <sub>2</sub> O Insol.	0.44 1.72 27.28 39.83 29.92 0.55	Fe" Al Fe" S H <sub>2</sub> O	0.08 0.41 4.18 6.09 20.34	Oxygen	0.08 0.62 6.27 18.27	25.24	
-	99.74					•	

The atomic content approaches  $X(OH)_2Fe^{\prime\prime\prime}_4(SO_4)_6$ .  $nH_2O$  where X is one oxygen equivalent of  $(Fe^{\prime\prime\prime}, Al)$  and n appears to be 19. A fuller discussion of this and other analyses of copiapite is reserved for the projected paper on the composition and optics of the mineral.

#### OBSERVATIONS ON CRYSTALS FROM SIERRA GORDA

Ungemach's extensive work on copiapite (1935) was done on complex crystals from Sierra Gorda. Adopting at first a setting with the plane of platy development and cleavage as the base ("orientation ancienne") he finally changed to a setting in which this plane is the side pinakoid ("notation définitive"). In this respect Ungemach's final setting resembles ours but otherwise it is

Figs. 1-6.—Crystals of copiapite from Chile. Fig. 1.—Chuquicamata; material used for 2-circle measurement, x-ray measurement, and analysis. Figs. 2, 3.—Sierra Gorda; two of Ungemach's crystals remeasured and redrawn in new position. Figs. 4, 5, 6.—Tierra Amarilla (pseudocopiapite); three of Ungemach's crystals remeasured and redrawn in new position. In Figs. 1, 4, 5 the positive end of the c-axis is directed upward and the crystals are viewed from the direction of b(010); in Figs. 2, 3, 6, representing crystals which are mainly developed at the negative end of the c-axis, this end is directed upward and the crystals are viewed from the direction of  $b(0\overline{10})$ .

radically different. Ungemach's preliminary setting is related to ours by the formula:

Ungemach (prelim.) to Authors:  $\frac{11}{26}0/001/\frac{11}{26}0$ Authors to Ungemach (prelim.):  $10\overline{1}/\overline{3}0\overline{3}/010$ 

The final setting of Ungemach stands in no less complicated relation to ours, since it defines a 12-fold cell in the structural lattice (Fig. 7):

Ungemach (final) to Authors:  $\frac{1}{2}0\frac{T}{6}/\frac{17}{22}0/\frac{T}{20}$ Authors to Ungemach (final):  $10\overline{1/121/303}$ 

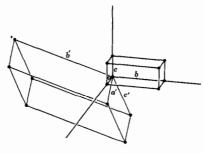


Fig. 7.—Crystal lattice of copiapite showing Ungemach's 12-fold morphological cell with  $a'[10\overline{1}]$ ,  $b'[\overline{12}1]$ ,  $c'[\overline{3}0\overline{3}]$ , in relation to the unit cell a, b, c, which conforms to the rules c < a < b, a and  $\beta$  obtuse.

The complexity of Ungemach's settings is indicated, as in othe similar cases, by the weakness of the chosen vertical zone. The axes of weak zones correspond to relatively long lattice periods and if such axes are taken as crystal axes the geometrical elements at likely to define multiple and therefore unsuitable lattice cells. On must admire Ungemach's consummate skill with the single-circ goniometer, but his critical remarks about the two-circle instrument (1935, p. 99) are the less effective when we note that the use of the instrument leads one always to choose the strongest zone ax (usually the shortest lattice period) as the vertical axis, and thus avoid the unfortunate settings that result from a feeble vertical zone.

The geometrical elements of Ungemach in his final setting:  $a:b:c=0.3010:1:0.7295;\ \alpha=99^{\circ}46', \beta=90^{\circ}30', \gamma=104^{\circ}21'$  give the following comparison with the elements of Palache and Peacock when transformed to our setting:

 $\begin{array}{cccc} a:b:c & \alpha & \beta \\ 0.4005:1:0.3971 & 93^{\circ}58\frac{1}{2}' & 102^{\circ}08' \\ 0.4058:1:0.4039 & 93^{\circ}50' & 102^{\circ}10' \\ 0.4037:1:0.4005 & 93^{\circ}51' & 101^{\circ}30' \end{array}$ 

Ungemach noted 146 forms on copiapit represented a number of the elaboratiskilful portrait drawings.

By combining Ungemach's geometri length of the vertical axis, c = 7.27 kX G = 2.134, both measured by Berry c Ungemach's collection, and using the resthe unit cell content, M = 1212.5, Unmaterial from Sierra Gorda may be expr the unit cell:

TABLE 3

COPIAPITE: ANALYSIS (UNGEMACH
Analysis Atoms

Analysis Atoms

Fe<sub>2</sub>O<sub>3</sub> 31.92 Fe''' 4.85
SO<sub>3</sub> 38.89 S 5.89
H<sub>2</sub>O [29.19] H<sub>2</sub>O 19.66

100.00

Again the cell content can be expressed as in which X is one oxygen equivalent of F

In view of the differences between t listed above, which correspond to anguexceeding half a degree, Peacock meast crystals which had kindly been made as reflections obtained were only fair, probedeterioration of the crystals brought about but it was possible to verify nearly all the on these particular crystals. The measu worth presenting since so many reflections somewhere between the angles calculated

<sup>&</sup>lt;sup>2</sup>In the warm dry atmosphere of the Mineral University whole drawers of salts from Chile hapowder with loss of water of crystallization.

#### Canadian Mineralogy

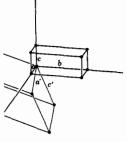
th's preliminary setting is related  $t_0$ 

.) to Authors:  $\frac{11}{26}0/001/\overline{\frac{11}{26}}0$  ach (prelim.):  $10\overline{1}/\overline{3}0\overline{3}/010$ 

nach stands in no less complicated s a 12-fold cell in the structural lattice

o Authors:  $\frac{1}{2}0\overline{0}$  ach (final):  $10\overline{1}$ 

 $\begin{array}{c} \frac{1}{2}0_{\overline{6}}^{\overline{1}}/\frac{\overline{1}}{2}0/\frac{\overline{1}}{2}0_{\overline{6}}^{\overline{1}} \\ 10\overline{1}/\overline{12}1/\overline{3}0\overline{3} \end{array}$ 



iapite showing Ungemach's 12-fold morpho $0\overline{3}$ ], in relation to the unit cell  $a,\ b,\ c$ , which and  $\beta$  obtuse.

h's settings is indicated, as in other sof the chosen vertical zone. The I to relatively long lattice periods and tal axes the geometrical elements are herefore unsuitable lattice cells. One summate skill with the single-circle narks about the two-circle instrument tive when we note that the use of this to choose the strongest zone axis priod) as the vertical axis, and thus to gs that result from a feeble vertical

of Ungemach in his final setting:  $\alpha = 99^{\circ}46'$ ,  $\beta = 90^{\circ}30'$ ,  $\gamma = 104^{\circ}21'$  n with the elements of Palache and our setting:

Ungemach noted 146 forms on copiapite from Sierra Gorda, and represented a number of the elaborately developed crystals by skilful portrait drawings.

By combining Ungemach's geometrical cell with the absolute length of the vertical axis, c = 7.27 kX, and the specific gravity, G = 2.134, both measured by Berry on a typical crystal from Ungemach's collection, and using the resulting molecular weight of the unit cell content, M = 1212.5, Ungemach's analysis of his material from Sierra Gorda may be expressed in terms of atoms in the unit cell:

TABLE 3
COPIAPITE: ANALYSIS (UNGEMACH) AND CELL CONTENT

Analysis			Atom			
Fe <sub>2</sub> O <sub>3</sub> SO <sub>3</sub> H <sub>2</sub> O	31.92 38.89 [29.19]	Fe''' S H <sub>2</sub> O	4.85 5.89 19.66	Oxygen	7.27) 17.67	24.94

100.00

Again the cell content can be expressed as  $X(OH)_2Fe'''_4(SO_4)_6.nH_2O$  in which X is one oxygen equivalent of Fe''' and n is apparently 19.

In view of the differences between the three sets of elements listed above, which correspond to angular differences sometimes exceeding half a degree, Peacock measured some of Ungemach's crystals which had kindly been made available by Donnay. The reflections obtained were only fair, probably due in part to some deterioration of the crystals brought about by atmospheric changes; but it was possible to verify nearly all the forms noted by Ungemach on these particular crystals. The measurements, which are hardly worth presenting since so many reflections were poor, often fall somewhere between the angles calculated from Palache's and Unge-

<sup>&</sup>lt;sup>2</sup>In the warm dry atmosphere of the Mineralogical Laboratories at Harvard University whole drawers of salts from Chile have unfortunately crumbled to powder with loss of water of crystallization.

mach's elements. It was not felt that these measurements should be given much weight in considering the question of the best geometrical elements for the mineral. Two of the remeasured crystals of Ungemach from Sierra Gorda are shown redrawn in our setting (Figs. 2, 3). Both crystals were more fully developed at the negative end of the c-axis which has therefore been directed upward while the broad face to the front is  $b(0\bar{1}0)$ .

#### OTHER OBSERVATIONS

In addition to normal copiapite, Ungemach distinguished an aberrant variety from Tierra Amarilla, Chile, which he named "pseudocopiapite." This material, which was said to have the same composition as normal copiapite, gave triclinic elements differing significantly from those of normal copiapite only in regard to the interaxial angles. Berry (1938b) remeasured fifteen crystals of Ungemach's pseudocopiapite and obtained elements which compare as follows with those of Ungemach transformed to our setting:

Comparing these values with the previously given elements for normal copiapite one sees that there is now a significant difference only in the interaxial angles a. This difference might be due to an undetected difference in chemical composition which is allowed by variation in the terms X and n of the proposed general formula.

It remains to note some observations on very imperfect crystals of a cuprian copiapite from Chuquicamata, Chile, named *cuprocopiapite* by Bandy (1938). The material consists of tiny packs of minute weakly cohering nearly parallel green plates with the typical rhombic outline of copiapite. The broad face has a pearly lustre and even when freshly cleaved gives only a blurred signal. With the cleavage polar, trains of feeble reflections were obtained from the edges of the rhombic plates. From twelve pairs of measurements the average interzonal angle is 77° 42′, which is close to the interzonal angle [001]: [100] of copiapite, for which Ungemach gives 77°  $48\frac{1}{2}$ ′, Palache 77° 50′. In one of these zones a fair reflection is frequently seen at the mean polar distance  $B = 84^{\circ}$  13′ as compared to Palache's  $B = 83^{\circ}$  59′ for (001). In the other zone several signals

gave the mean angle  $B=83^{\circ} 22'$  which the angle to (100) for which Palache §  $B=80^{\circ} 04\frac{1}{2}'$ . Thus the copper salt is the crystals do not permit a determination.

#### FORMAL CRYSTALLOGRA

Setting, elements, and angles. We sent the geometrical crystallography and graphical manner. The crystallo well established, since it conforms to t structural lattice in a widely used un which was accepted for copiapite by I

In regard to the choice of the geometrical elements, we have conside Palache's values, which agree best w analysed material from Chuquicama which accord best with his measureme Sierra Gorda; (3) average values, which results into account. The differences elements that have been derived for c to the quality of the crystals which g tions; probably of greater importance in chemical composition which will be projected paper on composition and c better to retain a set of values associa sition rather than to submerge real v Since Ungemach's measurements repre of crystals, and his analysis indic (X = Fe''') without monovalent or cided to adopt these values in our sett calculated angles (hkl): (010) also af corresponding B-angles.

Table 4 gives the adopted elements commoner forms of copiapite as indica statistics and our observations. In a circle angles,  $\varphi$ ,  $\rho$ , and the interfacia A, B, C, we give the interzonal ang

s not felt that these measurements should considering the question of the best geomineral. Two of the remeasured crystals a Gorda are shown redrawn in our setting als were more fully developed at the negath has therefore been directed upward while nt is  $b(0\overline{1}0)$ .

## HER OBSERVATIONS

al copiapite, Ungemach distinguished an Tierra Amarilla, Chile, which he named s material, which was said to have the same copiapite, gave triclinic elements differing of normal copiapite only in regard to the y (1938b) remeasured fifteen crystals of spite and obtained elements which compare Ungemach transformed to our setting:

98°59′ Ungemach 91°18½' 102°04' 98°50′ 102°22′ Berry 91°22′ s with the previously given elements for es that there is now a significant difference gles a. This difference might be due to an chemical composition which is allowed by  $\zeta$  and n of the proposed general formula. me observations on very imperfect crystals rom Chuquicamata, Chile, named cuprocoi). The material consists of tiny packs of nearly parallel green plates with the typical apite. The broad face has a pearly lustre cleaved gives only a blurred signal. With ns of feeble reflections were obtained from ic plates. From twelve pairs of measurezonal angle is 77° 42′, which is close to the 100] of copiapite, for which Ungemach gives '. In one of these zones a fair reflection is ean polar distance  $B=84^{\circ}\,13'$  as compared ' for (001). In the other zone several signals gave the mean angle  $B=83^{\circ}\,22'$  which is only roughly similar to the angle to (100) for which Palache gives  $B=79^{\circ}\,34'$ , Ungemach  $B=80^{\circ}\,04\frac{1}{2}'$ . Thus the copper salt is close to copiapite in form, but the crystals do not permit a determination of the elements.

#### FORMAL CRYSTALLOGRAPHIC PRESENTATION

Setting, elements, and angles. We are now in a position to present the geometrical crystallography of copiapite in formal tabular and graphical manner. The crystallographic setting is, we believe, well established, since it conforms to the properly chosen cell of the structural lattice in a widely used unique conventional orientation which was accepted for copiapite by Ungemach.

In regard to the choice of the numerical values for the geometrical elements, we have considered several alternatives: (1) Palache's values, which agree best with the measurements on the analysed material from Chuquicamata; (2) Ungemach's values, which accord best with his measurements on analyzed material from Sierra Gorda; (3) average values, which might take Peacock's x-ray results into account. The differences between the several sets of elements that have been derived for copianite may be due in part to the quality of the crystals which generally give only fair reflections; probably of greater importance is the considerable variation in chemical composition which will be further brought out in the projected paper on composition and optics. In that case it will be better to retain a set of values associated with a particular composition rather than to submerge real variation in a general average. Since Ungemach's measurements represent much the largest number of crystals, and his analysis indicates pure ferrian copiapite (X = Fe''') without monovalent or divalent bases, we have decided to adopt these values in our setting. In this way Ungemach's calculated angles (hkl): (010) also afforded a useful check on our corresponding B-angles.

Table 4 gives the adopted elements and calculated angles for the commoner forms of copiapite as indicated by Ungemach's frequency statistics and our observations. In addition to the standard two-circle angles,  $\varphi$ ,  $\rho$ , and the interfacial angles to the axial planes, A, B, C, we give the interzonal angle Z which, together with B,

## TABLE 4 COPIAPITE— $R(OH)_2Fe^{\prime\prime\prime}_4(SO_4)_620H_2O*$

Triclinic; pinakoidal—1

 $a:b:c=0.4005:1:0.3971; \ \alpha=93^{\circ}58\frac{1}{2}', \ \beta=102^{\circ}08', \ \gamma=98^{\circ}50'$   $p_{0}:q_{0}:r_{0}=1.0010:0.3929:1; \ \lambda=83^{\circ}58', \ \mu=77^{\circ}03\frac{1}{2}', \ \nu=80^{\circ}04\frac{1}{2}'$   $p_{0}'=1.0301, \ q_{0}'=0.4043, \ x_{0}'=0.2161, \ y_{0}'=0.1081$ 

E.	rms		ρ	A	В	С	<b>Z</b>
F0		φ					77°48½′
с	(001)		13°35′	77° 03½′			$\begin{array}{c} 77 & 48_{\overline{2}} \\ 0 & 00 \end{array}$
b	(010)	0 00	90 00		0 00	83 58	
а	(100)	$80\ 04\frac{1}{2}$	90 00	0 00	$80\ 04\frac{1}{2}$	77 $03\frac{1}{2}$	0 00
m	(110)	$60\ 10^{\frac{1}{2}}$	90 00	19 54	$60\ 10\frac{1}{2}$	$76\ 26\frac{1}{2}$	0 00
$M^{\frac{1}{2}}$	$(2\overline{1}0)$	$91\ 23\frac{1}{2}$	90 00	11 19	$91\ 23\frac{1}{2}$	$78\ 01\frac{1}{2}$	0 00
M	$(1\overline{1}0)$	102 36	90 00	$22\ 31\frac{1}{2}$	102 36	$79\ 30\frac{1}{2}$	0 00
M2	$(1\overline{2}0)$	121 53	90 00	$41\ 48\frac{1}{2}$	121 53	$82\ 56\frac{1}{2}$	0 00
M3	$(1\overline{3}0)$			55 30	$135\ 34\frac{1}{2}$	$85 \ 52\frac{1}{2}$	0 00
$d^{\frac{1}{2}}$	(012)	$34 \ 51\frac{1}{2}$		$75\ 34\frac{1}{2}$	73 08	10 50	
ı	(011)	22 52	$29\ 04\frac{1}{2}$	74 44½	63 24	20 34	77 48½
d	(011)	13 16		74 20			$77\ 48\frac{1}{2}$
d2	(021)	$9\ 17\frac{1}{2}$				-	
d3	(031)	9 172	99 1 <del>4</del>	14 422	01 102	-	-
e	$(0\overline{1}1)$	143 53		$81\ 15\frac{1}{2}$	106 09		77 $48\frac{1}{2}$
f	(101)	<b>7</b> 6 56	$51\ 38\frac{1}{2}$		$79\ 47\frac{1}{2}$		$39\ 05\frac{1}{2}$
$g\frac{1}{2}$	$(\overline{1}02)$	-86 12	$16\ 16\frac{1}{2}$	$105\ 47\frac{1}{2}$	88 56	$28\ 44\frac{1}{2}$	$106 \ 14\frac{1}{2}$
g	(101)	-94 58	38 43	$128 \ 32\frac{1}{2}$	$93\ 06\frac{1}{2}$	$51\ 29\frac{1}{2}$	128 361
g2		-97 45	$61\ 21\frac{1}{2}$	151 17	96 48	$74\ 13\frac{1}{2}$	151 08
$u^{\frac{3}{2}}$	$(\overline{13}2)$	$-153 \ 37$	33 14	$108\ 56\frac{1}{2}$	$119\ 24\frac{1}{2}$	$44\ 45$	$106 \ 14\frac{1}{2}$
<i>u</i> 2	(102)	200 0.		-			
<b>p</b> 3	(131)	39 24	$62\ 43\frac{1}{2}$	47 37		$50\ 29\frac{1}{2}$	$39\ 05\frac{1}{2}$
$r^{\frac{1}{2}}$	$(\overline{2}\overline{1}2)$	-10847	40 09				$128 \ 36\frac{1}{2}$
r	$(\overline{1}\overline{1}1)$	$-120\ 40\frac{1}{2}$	$42\ 52\frac{1}{2}$	129 31	$110 \ 18\frac{1}{2}$	56 26	$128\ 36\frac{1}{2}$
r2	$(\overline{12}1)$	$-137\ 42\frac{1}{2}$	49 53	127 11	124 27	$62\ 41\frac{1}{2}$	$128\ 36\frac{1}{2}$
		-67   15		123 26	<b>75 20</b>		$128\ 36\frac{1}{2}$
s s <b>2</b>		$-47  10^{\frac{1}{2}}$		$116\ 29\frac{1}{2}$	$59 \ 59\frac{1}{2}$	53 22	$128\ 36\frac{1}{2}$
J.	()	J <b></b> 2					151 051
w3	$(\overline{23}1)$	-128 50		$143 \ 32\frac{1}{2}$			
w6	$(\overline{26}1)$	$-145 50\frac{1}{2}$		131 39			151 0/2
v3	$(\overline{2}31)$	$-61\ 57\frac{1}{2}$	64 03	$135\ 08\frac{1}{2}$	$64 59 \frac{1}{2}$	$72\ 21\frac{1}{2}$	$151 \ 07\frac{1}{2}$

<sup>\*</sup>Composition indicated by a large number of analyses.

gives azimuth and distance from the r with Table 1 will show the differences according to the two sets of elements. how future measurements on analysed alternative values.

Form-list and projection. To conclography we give a complete list of the a gnomonic projection of the whole for the complicated relation between U and the fact that Ungemach's form-le Lévy'' notation in the discarded settin full list of forms in both notations. T tance of the forms we have marked w one of us while Ungemach's actual number after each of his symbols. From the n in Tables 1 and 4 the forms are classed in the formal angle-table (Table 4), les out distinguishing mark).

The form-list of copiapite presente and tables forms can be denoted simpl in drawings and projections briefer s letters, are convenient and generally t present a problem. In the case of we puted setting, such as calcite (Palache, to retaining and, if necessary, adding to letters, drawing on Roman, Greek, a case and capitals, and combining these as single, double, or triple dots, to pro distinct characters. But this procedur and errors even in the hands of careful relatively little known copiapite, on t tional set of form-letters, and the pre tically altered to a new one which we h remain unchanged. An appropriate se fore be used to denote the numerous cr

The forms of copiapite all lie on a sl b' (Fig. 8). The transverse zone [aa

TABLE 4
:—R(OH)<sub>2</sub>Fe'''<sub>4</sub>(SO<sub>4</sub>)<sub>6</sub>20H<sub>2</sub>O\*
riclinic; pinakoidal—\(\bar{1}\)

3971;  $\alpha = 93^{\circ}58\frac{1}{2}'$ ,  $\beta = 102^{\circ}08'$ ,  $\gamma = 98^{\circ}50'$ 9:1;  $\lambda = 83^{\circ}58'$ ,  $\mu = 77^{\circ}03\frac{1}{2}'$ ,  $\nu = 80^{\circ}04\frac{1}{2}'$ 0.4043,  $x_0' = 0.2161$ ,  $y_0' = 0.1081$ 

, -	′ •			
ρ	A	B	C .	<i>Z</i>
13°35′	77° 03½′	83°58′	0°00′	$77^{\circ}48\frac{1}{2}'$
90 00	$80\ 04\frac{1}{2}$	0 00	83 58	0 00
90 00	0.00	80 041	$77\ 03\frac{1}{2}$	
90 00				
90 00	19 54	$60\ 10^{\frac{1}{2}}$	$76\ 26\frac{1}{2}$	0 00
	11 10	01 93Ŧ	7× 111 *	0 00
90 00	$\frac{11}{22}$ $\frac{19}{31\frac{1}{2}}$	102 36	$79\ 30\frac{1}{2}$	0 00
90 00	$41\ 48\frac{1}{2}$	121 53	$82  ext{ } 56\frac{1}{2}$	0 00
90 00	55 30	$135\ 34\frac{1}{2}$	$85 \ 52\frac{1}{2}$	0 00
20 42½	$75 \ 34\frac{1}{2}$	<b>73</b> 08	82 56½ 85 52½ 10 50	77 $48\frac{1}{2}$
29 04 1	$74\ 44\frac{1}{2}$	63 24	$\begin{array}{c} 20 \ 34 \\ 35 \ 49\frac{1}{2} \\ 46 \ 12\frac{1}{2} \end{array}$	77 $48\frac{1}{2}$
43 17	74 20	$48\ 08\frac{1}{2}$	$35\ 49\frac{1}{2}$	77 $48\frac{1}{2}$
53 14	$74\ 42\frac{1}{2}$	$37\ 45\frac{1}{2}$	$46\ 12\frac{1}{2}$	$77\ 48\frac{1}{2}$
20.08	$81\ 15\frac{1}{2}$	106 09	<b>22</b> 11	$77\ 48\frac{1}{2}$
F1 201	38 28	79 47 1	38 35⅓	39 U5½
16 16±	$105 \ 47\frac{1}{2}$	88 56	$28\ 44\frac{1}{2}$	$106\ 14\frac{1}{2}$
38 43	$128 \ 32\frac{1}{2}$	$93\ 06\frac{1}{2}$	$51\ 29\frac{1}{2}$	$128 \ 36\frac{1}{2}$
61 211	151 17	96 48	$74\ 13\frac{1}{2}$	151 08
33 14	$108 \ 56\frac{1}{2}$	$119 24\frac{1}{2}$	44 45	$106 \ 14\frac{1}{2}$
62 431	47 37	46 37	50 29½ 53 38	$39\ 05\frac{1}{2}$
40 09	$129 \ 34\frac{1}{2}$	101 59	53 38	128 $36\frac{1}{2}$
42 52½	129 31	110 18	53 38 56 26	$128 \ 36\frac{1}{2}$
1, 022				
49 53	127 11	124 27	$62\ 41\frac{1}{2}$	$128 \ 36\frac{1}{2}$
40 531	123 26	<b>75</b> 20	$\begin{array}{cccccccccccccccccccccccccccccccccccc$	128 $36\frac{1}{2}$
47 25	116 29½	59 59	$\frac{1}{2}$ 53 22	$128 \ 36\frac{1}{2}$
11 20				
66 45	$143 \ 32\frac{1}{2}$	125 11	80 03	$151\ 07\frac{1}{2}$
<b>72</b> 48	131 39	142 14	84 44	151 072
64 03	135 08	64 59	$\frac{1}{2}$ 72 21 $\frac{1}{2}$	$151 \ 07\frac{1}{2}$
h 1		1 .		

by a large number of analyses.

gives azimuth and distance from the normal to (010). Comparison with Table 1 will show the differences between corresponding angles according to the two sets of elements. It will be interesting to see how future measurements on analysed materials compare with these alternative values.

Form-list and projection. To conclude the descriptive crystal-lography we give a complete list of the observed forms (Table 5) and a gnomonic projection of the whole form-system (Fig. 9). In view of the complicated relation between Ungemach's notation and ours, and the fact that Ungemach's form-letters represent a "Millerized Lévy" notation in the discarded setting, it will be useful to give the full list of forms in both notations. To indicate the relative importance of the forms we have marked with x the forms noted by any one of us while Ungemach's actual number of observations given in Tables 1 and 4 the forms are classed as common (\*\*) and entered in the formal angle-table (Table 4), less common (\*), and rare (without distinguishing mark).

The form-list of copiapite presented a special problem. In text and tables forms can be denoted simply by their Miller indices, but in drawings and projections briefer symbols, usually single italic letters, are convenient and generally used. Long form-lists always present a problem. In the case of well known species in an undisputed setting, such as calcite (Palache, 1943), there is no alternative to retaining and, if necessary, adding to the best current set of formletters, drawing on Roman, Greek, and German alphabets, lower case and capitals, and combining these with additional symbols such as single, double, or triple dots, to produce the necessary number of distinct characters. But this procedure leads to practical difficulties and errors even in the hands of careful and scholarly workers. The relatively little known copiapite, on the other hand, has no traditional set of form-letters, and the previous setting has been drastically altered to a new one which we hope will recommend itself and remain unchanged. An appropriate set of new symbols may therefore be used to denote the numerous crystal forms.

The forms of copiapite all lie on a sheaf of zones meeting in b and b' (Fig. 8). The transverse zone [aca'] cuts this sheaf into zone-

segments which originate at their intersections with the zone [aca'] and terminate at b and b'. Conventional letters are given to the unit forms of the zone-segments as tabulated below.

Zo	one-Segment	Unit Form	Zone-Segment	Unit Form	
(	$\begin{array}{c} 100) - (010) \\ 100) - (0\overline{1}0) \\ 100) - (0\overline{1}0) \\ 001) - (0\overline{1}0) \\ 001) - (0\overline{1}0) \\ 001) - (100) \\ 001) - (\overline{1}00) \\ 101) - (0\overline{1}0) \\ 101) - (0\overline{1}0) \\ \end{array}$	$m(110)$ $M(1\overline{1}0)$ $d(011)$ $e(0\overline{1}1)$ $f(101)$ $g(\overline{1}01)$ $p(111)$ $q(1\overline{1}1)$	$\begin{array}{c} (\overline{1}01)-(0\overline{1}0) \\ (\overline{1}01)-(0\overline{1}0) \\ (\overline{1}02)-(0\overline{1}0) \\ (\overline{1}02)-(0\overline{1}0) \\ (\overline{2}01)-(0\overline{1}0) \\ (\overline{2}01)-(0\overline{1}0) \\ (\overline{4}01)-(0\overline{1}0) \\ (\overline{4}01)-(0\overline{1}0) \\ (\overline{5}01)-(0\overline{1}0) \\ \end{array}$	$r(\overline{11}1) \\ s(\overline{1}11) \\ t(\overline{1}22) \\ u(\overline{12}2) \\ v(\overline{2}11) \\ w(\overline{21}1) \\ x(411) \\ y(4\overline{1}1) \\ z(\overline{5}11)$	

Any pole in a zone-segment, other than the end-poles and the units, are simply designated by the letter of the unit followed by a number which gives the gnomonic distance of the pole from the

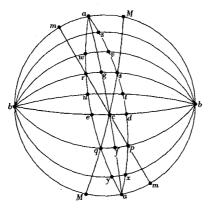
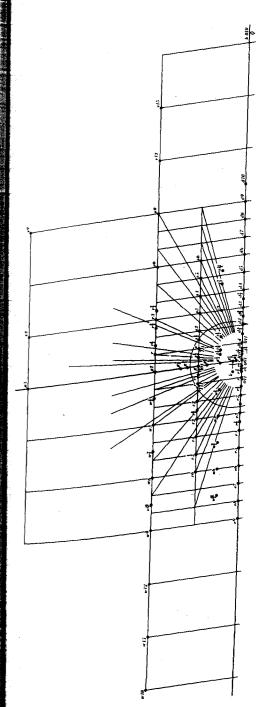


Fig. 8.—Stereographic projection of the principal zones of copiapite with the letters used to denote the forms in each zone-segment.

origin in terms of the unit gnomonic distance. For example, the pole (131) lies on the zone-segment f-b at a distance from the origin f which is three times the unit distance f-p. The form (131) therefore receives the symbol p3. Similarly (292) is lettered  $p\frac{9}{2}$ , (270)

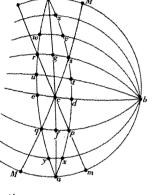


# S TO CANADIAN MINERALOGY

at their intersections with the zone  $[a_{ca'}]$  i. Conventional letters are given to the nents as tabulated below.

	The state of the	low.	-11 (0
Form	Zone-Segment	Unit Form	
110) )11) )11)	$\begin{array}{c} (\overline{1}01) - (0\overline{1}0) \\ (\overline{1}01) - (010) \\ (\overline{1}02) - (010) \\ (\overline{1}02) \end{array}$	$r(\overline{111})$ $s(\overline{111})$ $t(\overline{122})$	
101) 101) 111)	$\begin{array}{c} (\tilde{1} \bullet 2) - (0\tilde{1}0) \\ (\tilde{2}01) - (0\tilde{1}0) \\ (\tilde{2}01) - (0\tilde{1}0) \\ (\tilde{2}01) - (0\tilde{1}0) \\ (401) - (010) \end{array}$	$u(\overline{122}) \ v(\overline{211}) \ w(\overline{211})$	
.11)	$(\bar{5}01)$ — $(0\bar{1}0)$	$\begin{array}{c} x(411) \\ y(4\overline{1}1) \\ z(\overline{5}11) \end{array}$	į
men.			ı

gment, other than the end-poles and the ed by the letter of the unit followed by a gnomonic distance of the pole from the



ojection of the principal zones of comiapite with

t gnomonic distance. For example, the segment f—b at a distance from the origin  $\beta$ . The form (131) therese. Similarly (292) is lettered  $p_{\overline{z}}^9$ , (270)

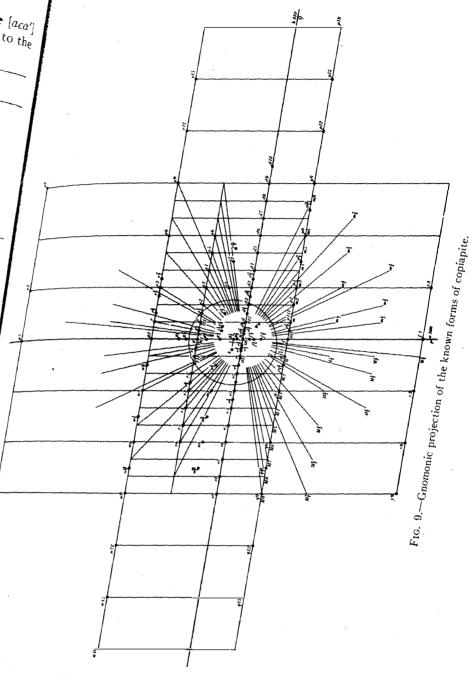


TABLE 5.—Con

A.	uthors	Obs.	Un	igemach	Obs.	At
$d_{\frac{2}{3}}^{2}$	(023)	_	$a\frac{1}{3}$	(319)	1	r
d	(011)**	x	<b>a</b> <sub>1</sub>	(113)	19	r2
$d^{\frac{3}{2}}$	(032)	x	$a_2$	(123)	1	$r\frac{5}{2}$
d2	(021)**	x	a <sub>3</sub>	(133)	13	$\parallel r\bar{3}$
$d_{\frac{5}{2}}^{5}$	(052)	x	αs	(143)	3	r4
d3	(031)**	x	αs	(153)	11	r5
$d^{\frac{7}{2}}$	(072)	- !	<b>a</b> 6	(163)	1	r6
$d\overline{4}$	(041)*	х	a 7	(173)	8	$s\frac{1}{2}$
d5	(051)*	x	ag	(193)	4	s
d6	(061)	x	<b>a</b> 11	(1.11.3)	2	$s\frac{3}{2}$
d7	(071)*	x	$a_{13}$	(1.13.3)	4	s2
d8	(081)	_	$a_{15}$	(1.15.3)	1	<i>s</i> 3
d9	(091)	-	a <sub>17</sub>	(1.17.3)	1	s4
d10	(0.10.1)	-	a 19	(1.19.3)	2	s5
$\frac{1}{3}$	(013)*	_	$\delta \frac{5}{3}$	$(3\overline{5}9)$	7	<i>s</i> 6
$e^{\frac{1}{2}}$	(012)*		$\delta_2$	$(1\overline{2}3)$	7	$w^{\frac{3}{2}}$
$\frac{2}{3}$	$(0\overline{2}3)*$	-	$\delta \frac{7}{3}$	$(3\overline{7}9)$	4	$\frac{\omega_2}{w3}$
2	$(0\overline{1}1)**$	x	$\delta_3$	$(1\overline{3}3)$	15	$w^{\frac{9}{2}}$
$\frac{3}{2}$	$(0\overline{3}2)*$		$\delta_4$	$(1\overline{4}3)$	5	$w_6$
$\tilde{2}$	$(0\overline{2}1)^*$	x	$\delta_5$	$(1\overline{5}3)$	8	$w^{\frac{15}{2}}$
<u>,5</u>	$(0\overline{5}2)$	_	$\delta_6$	$(1\overline{6}3)$	3	$w9^2$
3	$(0\overline{3}1)*$	x	$\delta_7$	(173)	4	w12
$\frac{7}{2}$	$(0\overline{7}2)$	_	$\delta_8$	(183)	2	w15
$\overline{4}$	$(0\overline{4}1)*$	_	$\delta_{\mathfrak{s}}$	$(1\overline{9}3)$	5	w13
25	$(0\overline{5}1)*$	_	$\delta_{11}$	$(1.\overline{11}.3)$	4	ข
6	$(0\overline{6}1)$	-	$\delta_{13}$	$(1.\overline{13.3})$ $(1.\overline{13.3})$	2	$v\frac{3}{2}$
7	(071)		$\delta_{15}$	$(1.\overline{15.3})$ $(1.\overline{15.3})$	3	$v\frac{5}{2}$
8	(081)		$\delta_{17}$	$(1.\overline{17.3})$	1	$v_{\overline{2}}$
9	$(0\bar{9}1)$	-	$\delta_{19}$	$(1.\overline{19.3})$	1	
7	(107)	_	$d^2$	$(1.19.3)$ $(1\overline{14})$		$v\frac{7}{2}$
į	(104)*	_	$d^{\frac{5}{2}}$		2	v6
<u>2</u>	(205)	-		$(1\overline{1}5) \\ (1\overline{1}7)$	7 2	υ9 10
	(101)**	,	$d_{\overline{2}}^{7}$		- 11	v12
4	(401)*	x	p .5	(001)	17	v15
1 8	$(\overline{1}08)$	-  -	$c_{\frac{5}{2}}$	$(\overline{1}15)$	4	<i>x</i> 3
3	$(\overline{105})*$	-	$d_{\frac{7}{6}}^{\frac{7}{6}}$	$(3\overline{3}7)$	3	у3
1 2	$(\overline{1}02)**$	- 1	$d^1$	$(1\overline{1}2)$	9	<b>y</b> 6
45	$(\overline{405})$	x	$d_1$	$(1\overline{1}1)$	18	<b>y</b> 9
	$(\overline{1}01)**$	-	$d^{1}_{\overline{6}}$	$(3\overline{3}1)$	1	z9
	(101)	<b>x</b> [	m	$(1\overline{1}0)$	21	z3

<sup>\*</sup>Common forms. \*Less common forms.

TABLE 5

			C	OPIAPITE	: OB	SERVE	ED FORMS	3			
Au	thors	Obs.	Ung	gemach	Obs.	Aut	hors	Obs.	Ung	gemach	Obs.
	(001)**	x	$\delta_1$	(113)	28	g <u>5</u>	(504)*	_	$c_{\frac{1}{6}}$	(331)	2
Ъ	(010)**	x	$g^1$	(010)	2	$g_{\overline{5}}^{\overline{7}}$	(705)	-	$c^{\frac{1}{4}}$	$(\overline{2}21)$	1
a	(100)**	x	$\gamma_1$	$(\overline{1}13)$	30	g2	$(\overline{2}01)**$	x	$c_1$	$(\overline{1}11)$	21
m8	(180)	- 1	$\gamma_{17}$	$(\bar{1}.17.3)$	1	g5	$(\overline{5}01)*$	-	$c^1$	$(\overline{1}12)$	8
<sub>m</sub> 7	(170)	x		$(\bar{1}.15.3)$	-	k	$(\overline{13}5)$	-	λ	$(1\overline{2}2)$	3
m6	(160)	x	$\gamma_{13}$	$(\bar{1}.13.3)$	2	h	(134)	-	$f_{\frac{5}{2}}^{5}$	(115)	1
m5	(150)	x	$\gamma_{11}$	$(\bar{1}.11.3)$	2	$t^{\frac{1}{2}}$	$(\overline{1}12)$	x		$(3\overline{1}3)$	_
m4	(140)	x	γ,	$(\overline{1}93)$	2	$t^{\frac{3}{4}}$	$(\overline{2}34)$	-	$o^1$	(101)	2
$m^{\frac{7}{2}}$	(270)	- {	$\gamma_8$	$(\overline{1}83)$	1	$t^{\frac{3}{2}}$	$(\overline{1}32)*$	x	$f_1$	(111)	8
$m\overline{3}$	(130)*	x	γ,	$(\overline{1}73)$	6	t3	$(\overline{1}62)^{*}$	-	$f_3$	(131)	5
$m^{\frac{5}{2}}$	(250)	-	$\gamma_6$	$(\overline{1}63)$	1	$t^{\frac{9}{2}}$	$(\overline{1}92)$	- (	$f_5$	(151)	2
$m^{2}$	(120)*	x	$\gamma_5$	$(\overline{1}53)$	6	$t^{\frac{2}{4}}$	$(\bar{2}.21.4)$	- [	$f_{6}$	(161)	1
$m^{\frac{3}{2}}$	(230)*	- ·	$\gamma_4$	$(\overline{1}43)$	4	$u^{\frac{3}{4}}$	$(\overline{2}\overline{3}4)^*$	-	$d_2$	$(1\overline{2}1)$	4
$m_{\frac{7}{3}}^{\frac{7}{4}}$	(340)	- 1	$\gamma^{\frac{1}{3}}$	$(\overline{3}.11.9)$	2	$u^{\frac{3}{2}}$	$(\overline{13}2)**$	-	$d_3$	$(1\bar{3}1)$	10
m	(110)**	x	$\gamma_3$	$(\overline{1}33)$	14	<i>u</i> 3	$(\overline{16}2)^{*}$	- [	$d_5$	$(1\overline{5}1)$	5
$m^{\frac{2}{3}}$	(320)	- 1	$\gamma \frac{7}{3}$	$(\bar{3}79)$	3	$u^{\frac{9}{2}}$	$(\overline{1}\overline{9}2)*$	-	$d_7$	$(1\overline{7}1)$	8
$m^{\frac{1}{2}}$	(210)*	-	$\gamma_2$	$(\overline{1}23)$	9	и6	$(\overline{1}.\overline{12}.2)$	-	$d_9$	$(1\overline{9}1)$	3
$m^{\frac{1}{3}}$	(310)	_	$\gamma \frac{5}{3}$	$(\bar{3}59)$	3	$u^{\frac{15}{2}}$	$(\overline{1}.\overline{15}.2)$	-	$d_{11}$	$(1.\overline{11}.1)$	3
$M^{\frac{1}{4}}$	$(4\overline{1}0)$	_	$\gamma^{\frac{1}{2}}$	$(\overline{2}16)$	3	$p^{\frac{1}{2}}$	(212)	x		(016)	-
$M^{\frac{1}{3}}$	(310)*	-	$\gamma \frac{1}{3}$	$(\bar{3}19)$	6	Þ	(111)	x		(013)	-
$M^{\frac{1}{2}}$	$(2\overline{1}0)**$	x	$a^3$	$(\overline{1}03)$	10	p2	(121)	x		(023)	_
$M_{\frac{2}{3}}$	$(3\overline{2}0)*$		$\beta_{\frac{1}{3}}$	$(\overline{31}9)$	5	<b>⊉</b> 3	(131)**	x ]	$i_1$	(011)	10
M	$(1\overline{1}0)**$	x	$\beta_1$	$(\overline{11}3)$	24	$p_{\frac{9}{2}}$	(292)	-	$i\frac{3}{2}$	(032)	1
$M_{\frac{4}{3}}^{4}$	$(3\bar{4}0)$	_	$\beta \frac{5}{3}$	$(\overline{35}9)$	1	<b>⊅</b> 6	(161)*	x	$i_2$	(021)	7
$M^{\frac{3}{2}}$	$(2\overline{3}0)*$	x	$\beta_2$	$(\overline{12}3)$	5	$p^{\frac{15}{2}}$	(2.15.2)	-	$i\frac{5}{2}$	(052)	2
$M^{2}_{2}$	$(1\overline{2}0)**$	x	$\beta_3$	$(\overline{13}3)$	10	<b>p</b> 9	(191)		$i_3$	(031)	3
$M_{\frac{5}{2}}^{\frac{5}{2}}$	$(2\overline{5}0)$	x	β4	$(\overline{14}3)$	2	p12	(1.12.1)	- 1	$i_4$	(041)	2
M3	$(1\overline{3}0)**$	х	$\beta_5$	(153)	10	<i>p</i> 15	(1.15.1)	-	$i_5$	(051)	2
$M_{\frac{7}{2}}$	(270)	х		$(\overline{163})$		<i>p</i> 18	(1.18.1)	-	$i_6$	(061)	2
$M_4^2$	$(1\overline{4}0)^*$	_	βτ	(173)	7	q	$(1\overline{1}1)$	-	$e^{\frac{1}{3}}$	$(0\overline{1}3)$	1
$M_{\frac{9}{2}}$	$(2\bar{9}0)$	-	$\beta_8$	$(\overline{18}3)$	1	$q^{\frac{3}{2}}$	$(2\overline{3}2)$	-	$e^{\frac{1}{2}}$	$(0\overline{1}2)$	1
$M_{5}^{2}$	$(1\overline{5}0)*$	_	$\beta_9$	(193)	8	<i>q</i> 3	(131)*	x	$e_1$	$(0\overline{1}1)$	9
<i>M</i> 6	$(1\overline{6}0)$	x	$\beta_{11}$	$(\overline{1}.\overline{11}.3$	) 2	<i>q</i> 6	$(1\overline{6}1)*$	x	$e_2$	$(0\overline{2}1)$	7
<i>M</i> 7	(170)	x	$\beta_{13}$	$(\overline{1}.\overline{1}\overline{3}.3$	) 2	$q^{\frac{15}{2}}$		_	$e_{2}^{5}$	$(0\overline{5}2)$	1
M8	(180)	٠.	$\beta_{15}$	$(\overline{1}.\overline{15}.3$		q9	$(1\overline{9}1)$	-	$e_3$	$(0\overline{3}1)$	. 2
M9	$(1\bar{9}0)$	_	$\beta_{17}$	$(\bar{1}.\bar{1}\bar{7}.3$	) 1	q12	$(1.\overline{12}.1)$		e4	$(0\overline{4}1)$	4
$d^{\frac{1}{3}}$	(013)*	_	$\delta \frac{1}{3}$	$(3\overline{1}9)$	6	q15	$(1.\overline{15}.1)$		e <sub>5</sub>	(051)	2
$d\frac{1}{2}$	(012)**	x	03	(103)	10	$r^{\frac{1}{2}}$	$(\overline{2}\overline{1}2)**$	-	$\frac{3}{2}g$	$(2\overline{3}0)$	10
2	/ .		1	` ,		11					

<sup>\*\*</sup>Common forms. \*Less common forms. Remaining forms rare.

TABLE 5
'E: Observed Forms

(	Obs.	Aut	hors	Obs.	Un	gemach	Obs
	28	g <u>5</u>	(504)*		$c_{\frac{1}{6}}$	$(\bar{3}31)$	2
	2	g 7/5	$(\bar{7}05)$	_	$c\frac{1}{4}$	$(\overline{2}21)$	1
	30	g2	$(\overline{2}01)**$	x	$c_1$	$(\overline{1}11)$	21
	1	g5	$(\overline{5}01)*$	-	$c^1$	$(\overline{1}12)$	8
	-	k	$(\overline{135})$	_	λ	$(1\overline{2}2)$	3
)	2	h	(134)	-	$f_{\frac{5}{2}}$	(115)	1
1	2	$t^{\frac{1}{2}}$	$(\overline{1}12)$	х		$(3\overline{1}3)$	
	2	$t^{\frac{3}{4}}$	$(\bar{2}34)$	-	$o^1$	(101)	2
	1	$t^{\frac{3}{2}}$	( <del>1</del> 32)*	x	$f_1$	(111)	8
	6	$t\bar{3}$	$(\bar{1}62)*$	-	$f_3$	(131)	5
	1	t 9/2	$(\overline{1}92)$	-	$f_5$	(151)	2
	6	$t^{\frac{2}{1}}$	$(\bar{2}.21.4)$	-	f 6	(161)	1
	4	$u^{\frac{3}{4}}$	$(\overline{23}4)*$	_	$d_2$	$(1\overline{2}1)$	4
	2	$u^{\frac{3}{2}}$	$(\overline{13}2)**$	-	$d_3$	$(1\overline{3}1)$	10
	14	u3	$(\overline{162})^*$	- 1	$d_5$	$(1\overline{5}1)$	5
	3	$u^{\frac{9}{2}}$	$(\bar{1}\bar{9}2)*$	-	$d_7$	$(1\overline{7}1)$	8
	9	<i>u</i> 6	$(\overline{1}.\overline{12}.2)$	_	$d_9$	$(1\bar{9}1)$	3
	3	$u^{\frac{15}{2}}$	$(\overline{1}.\overline{1}\overline{5}.2)$	_	$d_{11}$	(1.11.1)	
	3	$p^{\frac{1}{2}}$	(212)	x		(016)	_
	6	p	(111)	х		(013)	_
	10	<i>p</i> 2	(121)	х		(023)	_
	5	<i>p</i> 3	(131)**	х	$i_1$	(011)	10
	24	$p^{\frac{9}{2}}$	(292)	-	$i\frac{3}{2}$	(032)	1
	1	p6	(161)*	x	$i_2$	(021)	7
	5	$p^{\frac{15}{2}}$	(2.15.2)	-	$i\frac{5}{2}$	(052)	2
	10	<b>p</b> 9	(191)	-	$i_3$	(031)	3
	2	p12	(1.12.1)	-	$i_4$	(041)	2
	10	p15	(1.15.1)	_	$i_5$	(051)	2
	-	<i>p</i> 18	(1.18.1)	_	$i_6$	(061)	2
	7	q	$(1\overline{1}1)$	_	$e^{\frac{1}{3}}$	$(0\overline{1}3)$	1
	1	$q_{\frac{3}{2}}$	$(2\overline{3}2)$	_	$e^{\frac{1}{2}}$	$(0\overline{1}2)$	1
	8	<i>q</i> 3	$(1\overline{3}1)*$	x	$e_1$	$(0\overline{1}1)$	9
	2	96	$(1\overline{6}1)*$	x	$e_2$	$(0\overline{2}1)$	7
	2	$q^{\frac{1}{2}}$	$(2.\overline{15}.2)$	-	$e^{\frac{5}{2}}$	$(0\overline{5}2)$	1
	3	q9	$(1\bar{9}1)$	-	$e_3$	$(0\overline{3}1)$	2
	1	q12	$(1.\overline{12}.1)^{3}$	٠ _	e4	$(0\overline{4}1)$	4
	6	q15	$(1.\overline{15}.1)$	_	$e_{5}$	$(0\bar{5}1)$	2
	10	$r^{\frac{1}{2}}$	$(\overline{212})**$	_	$\frac{3}{2}g$	$(2\overline{3}0)$	10

mmon forms. Remaining forms rare.

TABLE 5.—Continued

, I	Authors		Ungemach		Obs.	Authors		Obs.	Ungemach		Obs.
$d_{\frac{2}{3}}^{2}$	(023)	_	$a^{\frac{1}{3}}$	(319)	1	r	(111)**	x	2g	(120)	17
d	(011)**	x	a1	(113)	19	r2	(121)**	_	3g	$(1\overline{3}0)$	11
$d^{\frac{3}{2}}$	(032)	х	a <sub>2</sub>	(123)	1	$r\frac{5}{2}$	$(\overline{252})$	_	$\frac{7}{2}g$	$(2\overline{7}0)$	1
$d\hat{2}$	(021)**	x	a <sub>3</sub>	(133)	13	$r\bar{3}$	(131)*	_ 1	4g	$(1\overline{4}0)$	7
$d^{\frac{5}{2}}$	(052)	х	a,	(143)	3	r4	141)*	- 1	5.g	$(1\overline{5}0)$	7
$d\tilde{3}$	(031)**	х	a <sub>5</sub>	(153)	11	r5	$(\overline{151})$	_ [	6g	$(1\overline{6}0)$	3
$d^{\frac{7}{2}}$	(072)	-	$a_6$	(163)	1	r6	$(\overline{161})$	_ ]	7g	(170)	1
d4	(041)*	x	<b>a</b> <sub>7</sub>	(173)	8	$S^{\frac{1}{2}}$	$(\overline{2}12)$	_	$_{2}h$	$(2\overline{1}0)$	3
d5	(051)*	x	a <sub>9</sub>	(193)	4	s	(111)**	x	$h^1$	(100)	20
d6	(061)	x	$a_{11}$	(1.11.3)	2	$s\frac{3}{2}$	( <del>2</del> 32)*	_	$h_2$	(210)	5
d7	(071)*	x	$a_{13}$	(1.13.3)	4	s2	(121)**	x	t	(110)	11
d8	(081)	- [	$a_{15}$	(1.15.3)	1	s3	(131)*	x	$g_2$	(120)	9
d9	(091)	-	$a_{17}$	(1.17.3)	1	s4	$(\bar{1}41)*$	_	g <sub>3</sub>	(130)	6
d10	(0.10.1)	-	<b>a</b> 19	(1.19.3)	2	<i>s</i> 5	$(\bar{1}51)*$	-	g <sub>4</sub>	(140)	4
$e^{\frac{1}{3}}$	$(0\overline{1}3)*$	-	$\delta \frac{5}{3}$	$(3\overline{5}9)$	7	<i>s</i> 6	$(\overline{1}61)$	- 1	g <sub>5</sub>	(150)	1
$e^{\frac{1}{2}}$	$(0\overline{1}2)^*$	-	$\delta_2$	$(1\overline{2}3)$	7	$w^{\frac{3}{2}}$	$(\overline{432})*$	-	$c_2$	$(\overline{1}21)$	4
$e^{\frac{2}{3}}$	$(0\overline{2}3)*$	-	$\delta \frac{7}{3}$	$(3\overline{7}9)$	4	$w\bar{3}$	$(\overline{231})**$	-	<i>C</i> <sub>3</sub>	$(\overline{1}31)$	10
e	$(0\overline{1}1)**$	x	$\delta_3$	$(1\overline{3}3)$	15	$w^{\frac{9}{2}}$	$(\overline{4}\overline{9}2)$	-	C4	$(\overline{1}41)$	2
$e^{\frac{3}{2}}$	$(0\overline{3}2)^*$	- [	$\delta_4$	$(1\overline{4}3)$	5	w6	$(\overline{261})**$	_	C <sub>5</sub>	$(\overline{1}51)$	11
e2	$(0\overline{2}1)^*$	x	$\delta_{\mathfrak{s}}$	$(1\overline{5}3)$	8	$w^{\frac{1}{2}}$	$(\overline{4}.\overline{15}.2)$	-	C <sub>6</sub>	$(\overline{1}61)$	1
$e^{\frac{5}{2}}$	$(0\overline{5}2)$	-	$\delta_{\scriptscriptstyle 6}$	$(1\overline{6}3)$	3	w9	$(\bar{291})*$	-	C7	$(\overline{1}71)$	6
e3	$(0\overline{3}1)*$	x	$\delta_7$	$(1\overline{7}3)$	4	w12	$(\overline{2}.\overline{12}.1)$	- [	C <sub>9</sub>	$(\overline{1}91)$	3
$e^{\frac{7}{2}}$	$(0\overline{7}2)$	-	$\delta_8$	$(1\overline{8}3)$	2	w15	$(\overline{2}.\overline{15}.1)$	- 1	$c_{11}$	$(\overline{1}.11.1)$	1
e4	$(0\overline{4}1)*$	-	$\delta_9$	$(1\overline{9}3)$	5	w18	$(\overline{2}.\overline{1}\overline{8}.1)$	-	$c_{13}$	$(\overline{1}.13.1)$	2
e5	$(0\overline{5}1)*$	-	$\delta_{11}$	$(1.\overline{1}\overline{1}.3)$	4	υ	$(\overline{2}11)$	x		$(\overline{3}13)$	_
e6	$(0\overline{6}1)$	-	$\delta_{13}$	$(1.\overline{13}.3)$	2	$v_{\overline{2}}$	$(\bar{4}32)*$	-	$a^1$	$(\overline{1}01)$	5
e <b>7</b>	$(0\overline{7}1)$	-	$\delta_{\scriptscriptstyle 15}$	$(1.\overline{15.3})$	3	$v_{\cdot 2}^{\underline{5}}$	$(\overline{4}52)$	x		$(\overline{323})$	_
e8	$(0\bar{8}1)$	-	$\delta_{17}$	$(1.\overline{17.3})$	1	v3	$(\overline{2}31)**$	x	$b_1$	$(\overline{1}\overline{1}1)$	9
ė9	$(0\bar{9}1)$	-	$\delta_{19}$	$(1.\overline{19}.3)$	1	$v^{\frac{7}{2}}$	$(\overline{4}72)$	x		$(\bar{3}\bar{4}\bar{3})$	_
$f^{\frac{1}{7}}$	(107)	-	$d^2$	$(1\overline{1}4)$	2	v6	$(\overline{2}61)*$	-	$b_3$	$(\overline{131})$	4
$f_{\frac{1}{4}}^{1}$	(104)*	-	$d^{\frac{5}{2}}$	$(1\overline{1}5)$	7	v9	$(\overline{2}91)*$	-	$b_5$	$(\overline{151})$	4
$f_{\overline{5}}^{\underline{2}}$	(205)	-	$d^{\frac{7}{2}}$	$(1\overline{1}7)$	2	v12	$(\overline{2}.12.1)$	-	$b_7$	$(\overline{1}\overline{7}1)$	1
f .	(101)**	x	p_	(001)	17	v15	$(\bar{2}.15.1)$	-	$b_9$	$(\overline{1}\overline{9}1)$	1
f4	(401)*	-	$c_{\frac{5}{2}}$	$(\overline{1}15)$	4	x3	(431)	-	θ	$(\bar{1}35)$	1
g <u>1</u>	$(\overline{1}08)$	-	$d\frac{7}{6}$	$(3\overline{3}7)$	3	у3	$(4\overline{3}1)$	-	$\beta \frac{5}{2}$	$(\overline{1}\overline{1}5)$	2
8 <del>1</del>	$(\bar{1}05)*$	-	$d^1$	$(1\bar{1}2)$	9	<b>y</b> 6	$(4\overline{6}1)$	-	$\eta_3$	$(\overline{135})$	1
g 1/2	$(\overline{1}02)**$	x	$d_1$	$(1\overline{1}1)$	18	<b>y</b> 9	$(4\bar{9}1)$	-	$\eta_5$	$(\overline{155})$	1
g <u>4</u>	$(\bar{4}05)$	-	$d\frac{1}{6}$	$(3\bar{3}1)$	1	<i>z</i> 9	$(\overline{5}31)$	-	$a^2$	$(\overline{1}02)$	1
g	( <del>101</del> )**	x	m	$(1\overline{1}0)$	21	z3	(591)	-	$\epsilon$	$(\overline{122})$	1

<sup>\*\*</sup>Common forms. \*Less common forms. Remaining forms rare.

becomes  $M_2^7$ , and so on. Only a few italic letters are used and the attached whole numbers or vulgar fractions are always written after the letter and on the same level. This gives a simple system which should present no typographic difficulty and could be adapted to other complex form-lists.

#### REFERENCES

Bandy, M. C. (1938): Mineralogy of three sulphate deposits of northern Chile-Am. Mineral., vol. 23, pp. 669-760.

BERRY, L. G. (1938a): A study of copiapite—M. A. Thesis, Univ. of Toronto (MS).

vol. 24, p. 182.

BERTRAND, E. (1881): De l'application du microscope à l'étude de la minéralogie—Bull. Soc. min. France, vol. 4, pp. 8-15.

DANA, E. S. (1892): System of mineralogy—New York.

Des Cloizeaux, A. (1881): Note sur les propriétés optiques de l'erytrozincite, de la raimondite et de la copiapite—Bull. Soc. min. France, vol. 4, pp.40-42

DOELTER, C. (1929): Handbuch der Mineralchemie, vol. 4(2)—Dresden & Leipzig.

GOLDSCHMIDT, V. (1913): Atlas der Krystallformen, vol. 2—Heidelberg.

HINTZE, C. (1930): Handbuch der Mineralogie, vol. 1, part 3(2)—Berlin & Leipzig.

LINCK, G. (1889): Beitrag zur Kenntniss der Sulfate von Tierra Amarilla bei Copiapo in Chile—Zeits. Kryst., vol. 15, pp. 1-28.

PALACHE, C. (1943): Calcite. An angle table and critical list—Contrib. Depl. Min. Pet. Harvard Univ., no. 259 (mimeog.).

Peacock, M. A. (1937): On normal triclinic face-symbols and the harmonic arithmetic rule (abs.)—Am. Mineral., vol. 22, p. 210.

vol. 24, p. 191. Crystallography of copiapite (abs.)—Am. Mineral.

Posnjak, E. & Merwin, H. E. (1922): The system, Fe<sub>2</sub>O<sub>3</sub>-SO<sub>3</sub>-H<sub>2</sub>O-Journ Am. Chem. Soc., vol. 44, pp. 1965-1994.

SCHARIZER, R. (1913): Beiträge zur Kenntnis der chemischen Constitution und der Genese der natürlichen Ferrisulfate VIII—Zeits. Kryst., vol. 52, pp. 372-398.

UNGEMACH, H. (1935): Sur certains minéraux sulfatés du Chili—Bull. Soc. frans Min., vol. 58, pp. 97-221. MINERALS FROM THE HIGH BEAVERDELL, BRI'

By A. B. Staples, Beave

H. V. WARREN, University of .

HE Highland-Bell Mine is now ■ Wallace Mountain, which is ne Mining Division, about 23 miles eas Columbia. The ore deposits of the by Reinecke (1915) who mentions, a gyrite, tetrahedrite, and native silv cussed the silver mineralization at 1 minerals and also argentite and po concerned mainly with the silver Mine, their modes of occurrence, out ties, and chemical compositions. paragenetic relations elsewhere. T R. B. Staples, Managing Director permission to publish this paper, a Vancouver, B.C., for the eleven ne which many weighed only 0.25 to 0

The ore-bearing veins of the High kettle Quartz Diorite (Jurassic), aro Monzonite. The veins range from thickness, with an average of about numerous faults. The silver miner tetrahedrite (freibergite), pyrargyrinative silver, all of which are descrabed given of the associated sphale pyrite are found in abundance but c widespread. Specular hematite, m been found by the authors but they ite (?) have been noted and fluorite

<sup>&</sup>lt;sup>1</sup>Descriptive mineralogy, abstracted by authors, from a longer manuscript.