



Supernova triggers for end-Devonian extinctions

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The Late Devonian was a protracted period of low speciation resulting in biodiversity decline, culminating in extinction events near the Devonian–Carboniferous boundary. Recent evidence indicates that the final extinction event may have coincided with a dramatic drop in stratospheric ozone, possibly due to a global temperature rise. Here we study an alternative possible cause for the postulated ozone drop: a nearby supernova explosion that could inflict damage by accelerating cosmic rays that can deliver ionizing radiation for up to ~ 100 ky. We therefore propose that the end-Devonian extinctions were triggered by supernova explosions at ~ 20 pc, somewhat beyond the “kill distance” that would have precipitated a full mass extinction. Such nearby supernovae are likely due to core collapses of massive stars; these are concentrated in the thin Galactic disk where the Sun resides. Detecting either of the long-lived radioisotopes ^{146}Sm or ^{244}Pu in one or more end-Devonian extinction strata would confirm a supernova origin, point to the core-collapse explosion of a massive star, and probe supernova nucleosynthesis. Other possible tests of the supernova hypothesis are discussed.

extinction | supernova | cosmic rays | ozone | isotope geology

The Late Devonian biodiversity crisis is characterized by a protracted decline in speciation rate occurring over millions of years (1, 2), punctuated by an extinction pulse (Kellwasser event) followed ~ 10 My later by a more moderate extinction (Hangenberg event) around the Devonian–Carboniferous boundary (DCB) ~ 359 My ago (3, 4). Marshall et al. (5) recently suggested that the Hangenberg event was associated with ozone depletion (see also ref. 6), in light of evidence such as malformations persisting in palynological assemblages on the order of many thousands of years. Ref. 7 argued that volcanic eruption and a large igneous province (LIP) triggered ozone depletion, whereas ref. 5 instead linked it to an episode of global warming not caused by LIP.

Previous work has not considered astrophysical sources of ionizing radiation, which are known to be possible causes of ozone depletion and concomitant ultraviolet-B (UV-B) increase that could trigger elevated extinction levels (see, e.g., ref. 8), as well as direct genetic damage. Here we consider whether astrophysical sources could account for the data in ref. 5, and whether any additional evidence could test for their occurrence.

The precise patterns prevalent during the DCB are complicated by several factors, including difficulties in stratigraphic correlation within and between marine and terrestrial settings and the overall paucity of plant remains (9). However, a general consensus seems to be emerging that there was first a loss of diversity in spores and pollen followed, after about 300 ky (10), by a pulse of extinctions of many plants including proto-trees, armored fish, trilobites, ammonites, conodonts, chitinozoans, and acritarchs, possibly coeval with the Hangenberg Crisis; this seems to have largely left intact sharks, bony fish, and tetrapods with five fingers and toes. The fact that these species disappeared

over multiple beds indicates that the extinction extended over at least thousands of years.

Refs. 5, 9, and 11 also report the discovery of spores from this episode with distinct morphologies including malformed spines and dark pigmented walls, features consistent with severely deteriorating environmental conditions, and UV-B damage following destruction of the ozone layer (11). However, more quantitative data are needed to study their variation during quiescent times in the fossil record.

Heating Mechanism for Ozone Depletion

Ref. 5 proposes an ozone depletion mechanism involving increased water vapor in the lower stratosphere caused by enhanced convection due to higher surface temperatures. Water vapor contributes to a catalytic cycle that converts inorganic chlorine (primarily HCl and ClONO₂) to free radical form (ClO). The ClO then participates in an ozone-destroying catalytic cycle. A similar set of cycles involving Br contributes to ozone depletion, but to a lesser extent (12). Increased ClO and decreased ozone following convective injection of water into the lower stratosphere has been verified by observation and modeling (12, 13). Ref. 5 argues that a period of exceptional and sustained warming would lead to the loss of the protective ozone layer via this mechanism.

This mechanism is important for lower stratosphere ozone depletion, and may have consequences for ground-level UV-B exposure (12). More detailed study is warranted. Until then, it is unclear whether this change would be sufficient to cause an extinction. There are several reasons for this.

First, the vertical extent of this ozone depletion mechanism should be limited to the lower stratosphere (~ 12 km to 18 km altitude) and does not overlap with the largest concentration of ozone, which occurs around 20 km to 30 km. So, while depletion may be significant in the lower stratosphere, the bulk of the ozone layer lies above this region and would not be affected. The total column density would be reduced, but not to the extent of a complete loss of the protective ozone layer.

Secondly, the duration of the effect should be relatively short, $\lesssim 1$ wk (12), since the injected water vapor is photolyzed and ClO is converted back to HCl and ClONO₂. Thus, unless convective transport of water vapor to the lower stratosphere, for example, by storms, is continuous (on week timescales), the ozone reduction will be episodic, not sustained. The effect is also seasonal, since strongly convective storms tend to be limited to the

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The authors declare no competing interest.

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spring/summer. While this is likely detrimental to surface life, most organisms have repair mechanisms that can cope with some short-duration UV-B exposure.

Thirdly, the effect is likely to be limited geographically, since strongly convective storms are not uniformly distributed and the enhanced water vapor is likely only to spread over ~ 100 km horizontally (12).

Finally, there is significant uncertainty as to the ozone depletion level needed to induce aberrations in pollen morphology and, even more critically, large-scale extinction. While the anthropogenic ozone “hole” over Antarctica has led to increased UV-B exposure, no crash in the ecosystem has resulted. This may partly be due to the seasonal nature of the change, as would be the case here as well. Recent work (14) has shown that short-term exposure to significant increases in UV-B does not result in large negative impacts on the primary productivity of ocean phytoplankton, and other organisms show a wide range of sensitivity (15, 16). The amount of column depletion over a given location in those cases was $\sim 50\%$. The depletion caused by the mechanism considered in ref. 5 seems unlikely to be that large. Hence, the convective transport of water vapor to the lower stratosphere may not be sufficient to induce a substantial extinction. It is thus worth considering other mechanisms for global ozone depletion.

Astrophysical Agents of Ozone Destruction and Biosphere Damage

Astrophysical mechanisms for biosphere damage include bolide impacts, solar proton events, supernova (SN) explosions, gamma-ray bursts, and neutron star mergers (kilonovae). Bolide impacts, gamma-ray bursts and solar proton events are essentially impulsive, and recovery of the ozone layer takes $\lesssim 10$ y (17), which is likely to avert lasting biosphere destruction. Moreover, these events and kilonovae are unlikely to recur frequently. Accordingly, we focus on SNe.

Supernovae (SNe) are prompt sources of ionizing photons: extreme UV, X-rays, and gamma rays. Over longer timescales, the blast collides with surrounding gas, forming a shock that drives particle acceleration. In this way, SNe produce cosmic rays, that is, atomic nuclei accelerated to high energies. These charged particles are magnetically confined inside the SN remnant, and are expected to bathe Earth for ~ 100 ky.

The cosmic ray intensity would be high enough to deplete the ozone layer and induce UV-B damage for thousands of years (18–21). In contrast to the episodic, seasonal, and geographically limited ozone depletion expected from enhanced convection, ozone depletion following an SN is long lived and global (see, e.g., refs. 16, 20, and 21) and is therefore much more likely to lead to an extinction event, even given uncertainties around the level of depletion necessary. [We note that, as well as the induced UV-B damage, cosmic rays could also cause radiation damage via muons produced when they impact the atmosphere (22).] The SN blast itself is unlikely to wreak significant damage on the biosphere, but may deposit detectable long-lived nuclear isotopes that could provide distinctive signatures, as we discuss later.

There are two main types of SNe: 1) massive stars ($\gtrsim 8M_{\odot}$) that explode as core-collapse SNe (CCSNe) and 2) white dwarfs that accrete from binary companions and explode as Type Ia SNe. These SN types have similar explosion energies, and both produce ionizing radiation able to damage the biosphere. However, their different nucleosynthesis outputs lead to different radioisotope signatures.

Near-Earth CCSNe are more likely than Type Ia SNe. We estimate the nearby CCSN frequency using a Galactic rate $\mathcal{R}_{\text{CCSN}} = (30 \text{ y})^{-1}$ and placing the Sun at a radius $R_{\odot} = 8.7$ kpc in a thin disk of scale radius 2.9 kpc and height 0.1 kpc (23). This gives a CCSN rate $\mathcal{R}_{\text{SN}} = e^{-R_{\odot}/R_0} r^3/3R_0^2 h_0 \approx 4 r_{20}^3 \text{ Gy}^{-1}$ within

$r_{20} = r/20$ pc from Earth. Hence a CCSN at a distance ≈ 2 times the “kill radius” of 10 pc is a plausible origin of the end-Devonian event(s). In contrast, the Type Ia SN rate is an order of magnitude smaller, as these events are spread over the ≈ 8 times larger volume of the thick disk.

Massive stars are usually born in clusters (OB associations), and are usually in binaries with other massive stars. Thus, if one CCSN occurred near the DCB, it is likely there were others. This could explain the Kellwasser and other enigmatic Devonian events, in addition to the Hangenberg event.

Possible Radioisotope Signatures of SNe

A CCSN close enough to cause a significant extinction would also deliver SN debris to the Earth as dust grains—micron- or submicron-sized particles created early after the explosion. Grains in the explosion would decouple from the plasma (gas) and propagate in the magnetized SN remnant until they were stopped or destroyed by sputtering during collisions (24).

The portion that reaches Earth would deposit in the atmosphere live (undecayed) radioactive isotopes. There is very little preexisting background for radioisotopes whose lifetimes are much shorter than the age of Earth. Those with lifetimes comparable to the time since the event would provide suitable signatures. The discoveries of live ^{60}Fe in the deep ocean, the lunar regolith, and Antarctic snow provide one such signal, which is interpreted as due to at least one recent nearby CCSN 2 My to 3 My ago at a distance of ~ 50 pc to 100 pc, which is compatible with the rate estimate given above (24).

Possible relic SN radioisotopes from the end-Devonian period with an age 360 My include ^{146}Sm (half-life 103 My), ^{235}U (half-life 704 My) and ^{244}Pu (half-life 80.0 My). The most promising signature may be provided by ^{244}Pu , which has also been discovered in deep-ocean crust and sediment samples deposited over the last 25 My (25). Moreover, it is absorbed into bones and retained during life (26), whereas uranium is absorbed during fossilization (27) and ^{146}Sm is soluble. There is a significant ^{235}U background surviving from before the formation of the solar system, with $(^{235}\text{U}/^{238}\text{U})_{\oplus} = 0.721 \pm 0.001\%$, so a significant detection above this background requires deposition attaining $^{235}\text{U}_{\text{SN}}/^{238}\text{U}_{\oplus} \gtrsim 3 \times 10^{-5}$. U-Pb dating has been used to date the end-Devonian extinction, with an uncertainty in the $^{235}\text{U}/^{238}\text{U}$ ratio that is much larger than this target sensitivity, but even a few atoms of nonanthropogenic ^{244}Pu in end-Devonian fossils would be unambiguous evidence for the r process in SNe.

We have estimated the terrestrial deposition of ^{146}Sm , ^{235}U , and ^{244}Pu by a nearby SN. The ^{146}Sm is a proton-rich (“ p process”) nucleus that might be produced by CCSNe or Type Ia SNe (28). Models for the p process (28) give $^{146}\text{Sm}/^{144}\text{Sm} \approx 0.01 - 2.5$, with the predicted core-collapse abundance typically around 0.2. Assuming a CCSN that produced a solar $^{144}\text{Sm}/^{16}\text{O}$ ratio, and ejected $M_{\text{ej}}(^{16}\text{O}) = 2M_{\odot}$, we estimate a total yield of ^{146}Sm in the ejecta of $\mathcal{N}(^{146}\text{Sm}) \approx 1.6 \times 10^{47}$ atoms. On the other hand, ^{244}Pu and ^{235}U are neutron-rich nuclei that are made by the rapid capture of neutrons, the r process, whose astrophysical sites are uncertain. There is evidence that kilonovae make at least some of the lighter r -process nuclei (29), but it is uncertain whether these events make the heavier nuclei of interest here. Assuming that CCSNe are the dominant r -process sites, we estimate yields of $\mathcal{N}(^{235}\text{U}, ^{244}\text{Pu}) \approx (3, 1.6) \times 10^{47}$ atoms per explosion.

The journey of SN-produced radioisotopes from explosion to seafloor is complex. Ejecta in dust most readily reaches Earth (30). The fraction of atoms in dust, f_{dust} , should be high for the refractory species of interest. Due to their high speeds, SN dust grains will easily overcome the solar wind and reach Earth (31). The fallout on Earth favors deposition at midlatitudes; additional dispersion occurs due to ocean currents (31). The

global average surface density of isotope i with half-life $t_{1/2}$ is $N_i = f_{\text{dust}} \mathcal{N}_{\text{ej},i} 2^{-t/t_{1/2}} / 16\pi r^2$ (30), with t as the time since the explosion. We thus find global-averaged end-Devonian surface densities of SN material,

$$N(^{146}\text{Sm}, ^{235}\text{U}, ^{244}\text{Pu}) \approx f_{\text{dust}} (1, 9, 0.3) \times 10^5 \text{ atoms/cm}^2 r_{20}^{-2}$$

after including the decay factors for each species. Unfortunately, this estimate implies a ratio of SN-produced ^{235}U to the background level in Earth's crust of $\mathcal{O}(10^{-10})$, which is undetectably small. On the other hand, there is no natural background to the prospective ^{244}Pu signal, which may be detectable in fossiliferous material. Its detectability depends on the temporal resolution of the available geological sample, whereas the possible detectability of the prospective ^{146}Sm signal depends also on the degree of dilution due to its solubility. Finally, if more than one SN occurred before the DCB, then each of these could deposit radioisotope signals.

Other Tests for SNe

Some hundreds or thousands of years after the optical and ionizing outburst, the cosmic ray and dust bombardment of Earth would begin, with several possible effects.

Cosmic ray ionization of the atmosphere and accompanying electron cascades may lead to more frequent lightning, increased nitrate deposition, and wildfires (32). The increased nitrate flux

might have led to CO_2 drawdown via its fertilization effect (33), thereby cooling the climate. There is evidence for cooling during the first stage of the DCB, although this occurred an estimated 300 ky before the radiation damage attested by the data on pollen and spores (5). Any increases in soot and carbon deposits during the end-Devonian could have been generated by increases in wildfires (32).

Cosmic rays striking the atmosphere produce energetic muons that can penetrate matter to a much larger depth than UV-B radiation. The radiation dose due to muons at Earth's surface (34) and in the oceans at depths of ≤ 1 km (22) could exceed, for many years, the current total radiation dose at Earth's surface from all sources. Therefore, in addition to comparing the effects of muons and UV-B radiation at or near the surface, they could be considered in end-Devonian extinctions of megafauna living at depth.

Finally, if there was one CCSN at the DCB, there may have been more, which may have been responsible for the Kellwasser and additional events. These could show evidence for ozone depletion and the other signatures above.

Data Availability. All study data are included in the article.

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1. A. Stigall, Speciation collapse and invasive species dynamics during the Late Devonian "mass extinction". *GSA Today* **22**, 4–9 (2012).
2. J. X. Fan *et al.*, A high-resolution summary of Cambrian to Early Triassic marine invertebrate biodiversity. *Science* **367**, 272–277 (2020).
3. S. Kaiser, M. Aretz, R. Becker, The global Hangenberg Crisis (Devonian–Carboniferous transition): Review of a first-order mass extinction. *Geol. Soc. Spec. Publ.* **423**, 387–437 (2016).
4. D. Bonda, S. Grasby, On the causes of mass extinctions. *Palaeogeogr. Palaeoclimatol. Palaeoecol.* **478**, 3–29 (2017).
5. J. Marshall, J. Lakin, I. Troth, S. Wallace-Johnson, UV-B radiation was the Devonian–Carboniferous boundary terrestrial extinction kill mechanism. *Sci. Adv.* **6**, eaba0768 (2020).
6. C. S. Cockell, Crises and extinction in the fossil record—a role for ultraviolet radiation? *Paleobiology* **25**, 212–225 (1999).
7. G. Racki, A volcanic scenario for the Frasnian Famennian major biotic crisis and other Late Devonian global changes More answers than questions. *Glob. Planet. Change* **189**, 103174 (2020).
8. A. L. Melott, B. C. Thomas, Astrophysical ionizing radiation and Earth: A brief review and census of intermittent intense sources. *Astrobiology* **11**, 343–361 (2011).
9. C. Prestianni, M. Sautois, J. Denayer, Disrupted continental environments around the Devonian–Carboniferous Boundary: Introduction of the tener event. *Geol. Belg.* **19**, 135–145 (2016).
10. P. Myrow *et al.*, High-precision U–Pb age and duration of the Latest Devonian (Famennian) Hangenberg Event, and its implications. *Terra. Nova* **26**, 222–229 (2014).
11. P. Filipiak, G. Racki, Proliferation of abnormal palynoflora during the end-Devonian biotic crisis. *Geol. Q.* **54**, 1–14 (2010).
12. J. G. Anderson, D. M. Wilmouth, J. B. Smith, D. S. Sayres, UV dosage levels in summer: Increased risk of ozone loss from convectively injected water vapour. *Science* **337**, 835–839 (2012).
13. J. G. Anderson *et al.*, Stratospheric ozone over the United States in summer linked to observations of convection and temperature via chlorine and bromine catalysis. *Proc. Natl. Acad. Sci. U.S.A.* **114**, E4905–E4913 (2017).
14. P. J. Neale, B. C. Thomas, Solar irradiance changes and phytoplankton productivity in Earth's ocean following astrophysical ionizing radiation events. *Astrobiology* **16**, 245–258 (2016).
15. B. C. Thomas, P. J. Neale, B. R. Snyder, Solar irradiance changes and photobiological effects at Earth's surface following astrophysical ionizing radiation. *Astrobiology* **15**, 207–220 (2015).
16. B. C. Thomas, Photobiological effects at Earth's surface following a 50 pc supernova. *Astrobiology* **18**, 481–490 (2018).
17. B. C. Thomas *et al.*, Gamma-ray bursts and the Earth: Exploration of atmospheric, biological, climatic, and biogeochemical effects. *Astrophys. J.* **634**, 509–533 (2005).
18. M. A. Ruderman, Possible consequences of nearby supernova explosions for atmospheric ozone and terrestrial life. *Science* **184**, 1079–1081 (1974).
19. J. Ellis, D. Schramm, Could a nearby supernova explosion have caused a mass extinction? *Proc. Natl. Acad. Sci. U.S.A.* **92**, 235–238 (1995).
20. N. Gehrels *et al.*, Ozone depletion from nearby supernovae. *Astrophys. J.* **585**, 1169–1176 (2003).
21. A. L. Melott, B. C. Thomas, M. Kachelrieß, D. V. Semikoz, A. C. Overholt, A supernova at 50 pc: Effects on the Earth's atmosphere and biota. *Astrophys. J.* **840**, 105 (2017).
22. A. L. Adrian, F. Marinho, L. Paulucci, Muon radiation dose and marine megafaunal extinction at the end-Pliocene supernova. *Astrobiology* **19**, 825–830 (2019).
23. S. M. Adams *et al.*, Observing the next galactic supernova. *Astrophys. J.* **778**, 164 (2013).
24. B. Fields *et al.*, Near-Earth supernova explosions: Evidence, implications, and opportunities. *Bull. Am. Astron. Soc.* **51**, 410 (2019).
25. A. Wallner *et al.*, Abundance of live ^{244}Pu in deep-sea reservoirs on Earth points to rarity of actinide nucleosynthesis. *Nat. Commun.* **6**, 5956 (2015).
26. Y. Takizawa, Plutonium in Japanese tissues. *J. Radiat. Res.* **23**, 198–203 (1982).
27. S. L. Koul, Uranium in fossil bones. *Radiat. Eff.* **43**, 7–11 (1979).
28. M. Arnould, S. Goriely, The p-process of stellar nucleosynthesis: Astrophysics and nuclear physics status. *Phys. Rep.* **384**, 1–84 (2003).
29. B. P. Abbott *et al.*, Multi-messenger observations of a binary neutron star merger. *Astrophys. J. Lett.* **848**, L12 (2017).
30. B. J. Fry, B. D. Fields, J. R. Ellis, Astrophysical shrapnel: Discriminating among near-Earth stellar explosion sources of live radioactive isotopes. *Astrophys. J.* **800**, 71 (2015).
31. B. J. Fry, B. D. Fields, J. R. Ellis, Radioactive iron rain: Transporting ^{60}Fe in supernova dust to the ocean floor. *Astrophys. J.* **827**, 48 (2016).
32. A. L. Melott, B. C. Thomas, From cosmic explosions to terrestrial fires? *J. Geol.* **127**, 475–481 (2019).
33. A. L. Melott, B. C. Thomas, B. D. Fields, Climate change via CO_2 drawdown from astrophysically initiated atmospheric ionization? *Int. J. Astrobiol.* **19**, 1–4 (2020).
34. B. C. Thomas *et al.*, Terrestrial effects of nearby supernovae in the early Pleistocene. *Astrophys. J. Lett.* **826**, L3 (2016).